In turn, the consistent uplift of the source areas, which at that time were the Silesian and Baška-Inwałd ridges surrounding the proto-Silesian basin, led to the sedimentation of coarse-movement formations in the form of olistoliths, marl gravels, slides of limestone blocks of the Štramberk type (Nowak, 1973; Słomka, 1986b, 2001). This would not have been possible if not for the intense synsedimentary tectonic movements in this part of the Tethys associated with the Neo-Cimmerian movements of the Alpine orogeny (Krobicki, 1996a, Golonka *et al*., 2003; cf. Middleton & Hampton, 1973; Stampfli *et al*., 1998; Shanmugam, 2002; Dasgupta, 2003; Payros & Pujalte, 2008 with literature cited there).

The Beriasian episode of resedimentation in the Proto-Silesian basin was almost exclusively associated with the destruction of carbonate platforms and uplifted calcareous flysch formations of the Cieszyn Limestone Formation. In the Valanginian, the bedrock erosion occurred (the first exotics of igneous and metamorphic rocks) and the increasing amount of clastic material led to the replacement of calcareous flysch sedimentation by silicoclastic flysch (Słomka, 2001).

The limestone flysch of the Late Jurassic/Early Cretaceous of the Cieszyn Limestone Formation finds good counterparts in many flysch/flysch-like systems in Europe and beyond (Payros & Pujalte, 2008 with literature cited there) and it is striking that, as in the Carpathians (Matyszkiewicz & Słomka, 1994), in many cases, one of the basic grain components of these formations are ooids, as resedimented grains from shallow-sea carbonate platforms, which are the source areas for calcareous turbidite/fluxoturbidite systems (Price, 1977; Ruiz-Ortiz, 1983; Wright & Wilson, 1984; Eberli, 1987; Cooper, 1989, 1990; Zempolich & Erba, 1999; Wright, 2004; Robin *et al*., 2005; Brookfield *et al*., 2006). On the other hand, within the resedimented formations, fragments (as olistoliths/exotics) of organic structures such as coral-algal reefs or microbial-sponge mud mounds can be found (Karpaty – Matyszkiewicz & Słomka, 2004, other locations – Payros & Pujalte, 2008), which document a wide range of shallow-water facies of carbonate platforms in source areas.

Taking into account all the threads mentioned – from the reconstruction of shallow-water carbonate sedimentation environments in the source areas, through the synsedimentary Neo-Cimmerian tectonics catalysing the formation of various resedimented formations, to the isochronous beginnings of submarine volcanic activity (teschenite – Grabowski *et al*., 2003, 2004), they should be associated with the increasing geotectonic activity of the Silesian Cordillera (Poprawa *et al*., 2004, 2005). Originally, it was probably a fragment of the East European platform, detached from it as a result of a short-term, abortive rift, which probably did not lead to full oceanization of the Proto-Silesian basin (Narębski, 1990) (Fig. 54), despite its inclusion in a continuous rift system between the opening Atlantic in the west, through the Ligurian – in the center to the Chornohora-Sinaia branch of the Ukrainian-Romanian Carpathians in the east (Mišík, 1992; Golonka & Krobicki, 2004; Krobicki *et al*., 2005; Rogoziński & Krobicki, 2006 with the literature cited there).

Stop 18 – Krakow vicinity – Middle-Upper Jurassic and Upper Cretaceous deposits

In Krakow and its vicinity the Mesozoic rocks of the North European platform are exposed. The platform is dissected by numerous faults into several horsts and grabens. The grabens are filled with the Miocene Molasse deposits, while horsts elevate Jurassic (mainly Oxfordian) limestones which sometime are covered by the Upper Cretaceous deposits of the Peri-Tethys realm of the northern margin of the Tethys Ocean.

Outcrops? Rocks? Palaeoenvironments? – a surprise for you… – thanks for coming, interesting conversations and company. See you next year in Krakow again on the 36th HKT (Himalaya–Karakorum–Tibet) Workshop at AGH University of Krakow – Welcome…!

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