## Andrzej Kostecki

The Oil and Gas Institute, Krakow

# Algorithm MG(F-K) of migration in model TTI anisotropy 

## Introduction

One of the most frequently encountered models in anisotropic medium is the model of Transverse Isotropy (TI). If the symmetry axis of model TI coincides with the vertical axis of right-angle coordinates, then we receive horizontal thin-layered system proposed by G. Postma [12] and called VTI (Vertical Transversely Isotropic). The application of algorithms of propagation and migration of ,isotropic" waves in anisotropic medium VTI results in deformations and relocation of reproduced structures $[6,10]$, and the higher the
anisotropic parameters are, the greater their extent is. (parameters by L. Thomsen [14]).

In the domain of wavenumbers and frequencies, the algorithm of migration $\mathrm{MG}(\mathrm{F}-\mathrm{K})$ was presented in anisotropic medium type VTI [11]. This algorithm uses approximative version of vertical wavenumber [5] with reference to medium VTI. This article presents the algorithm of migration $\mathrm{MG}(\mathrm{F}-\mathrm{K})$ in monoclinal medium marked as TTI (Tilted Transversely Isotropic) model, whose symmetry axis is tilted at $\theta$ angle to the vertical axis (Fig. 1).

## Algorithmic solutions

In case of a thin-layered arrangement, arbitrarily oriented in relation to the Cartesian coordinate system $x, y, z$, it is appropriate to use general law of tensor rotation in Bond`s formulation [2, 3, 13] which allows to obtain relation between matrix $D^{\varphi \theta}$ of elastic modules in measuring coordinate system $x, y, z$, and appropriate matrix $C$ of these modules in coordinate system $x^{\prime}, y^{\prime}, z^{\prime}$.

The is the ensuing relation:

$$
\begin{equation*}
D^{\varphi \theta}=R_{(\varphi \theta)} C R_{(\varphi \theta)}^{T} \tag{1}
\end{equation*}
$$

where $\varphi$ denotes the angle of rotation of system $x^{\prime}, y^{\prime}, z^{\prime}$ with regard to axis $z$, whereas $\theta$ is the tilt angle of the symmetry plane of isotropy TI.

The dimensions of matrices $R_{\varphi \theta}$ and $R_{\varphi \theta}^{T}$ are $6 \times 6$ and they transform the matrix of elastic modules $C$ into symmetric matrix $D^{\varphi \theta}$. Matrices $R_{\varphi \theta}$ and $R_{\varphi \theta}^{T}$ transpose the vectors of stress and strain from system $x^{\prime}, y^{\prime}, z^{\prime}$ to system $x, y, z$ by means of rotational matrix:

$$
A=\left|\begin{array}{ccc}
\cos \theta \sin \varphi & \cos \varphi & \sin \theta \sin \varphi  \tag{2}\\
-\cos \theta \cos \varphi & \sin \varphi & -\sin \theta \cos \varphi \\
-\sin \theta & 0 & \cos \theta
\end{array}\right|
$$



Fig. 1. Geometrical model of monoclinal thin-layered system, angle $\theta$ is the tilt angle between axis x ' and horizontal plane. Rotation angle $\varphi$ of system $\mathrm{x}^{\prime}, \mathrm{y}^{\prime}, \mathrm{z}$ ' towards the axis is $90^{\circ}$

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The subject of discussion will be the situation when the rotation angle $\varphi=90^{\circ}$, i.e. when the coincidence of axis $y$ and $y^{\prime}$ occurs (Fig. 1). Then, similarly as for the medium VTI [4] the shear waves of type SH may be separated from longitudinal waves P and shear waves SV. It means that displacements $U_{y}$ of oscillating particles of the medium towards axis y are independent from displacements $U_{x}$ and $U_{z}$ in directions $x$ and $z$ respectively. Thus, only components $U_{x}$ and $U_{z}$ can be discussed, assuming that $U_{y}$ and its derivatives equal zero.

Symmetrical matrix $C$, dimensions $6 \times 6$, represents components of tensor $C_{i j l k}$ in the medium of transverse isotropy TI. In abbreviated notation by Voigt [14] this matrix can be presented in this way:

$$
C=\left|\begin{array}{cccccc}
C_{11} & C_{12} & C_{13} & 0 & 0 & 0  \tag{3}\\
C_{12} & C_{11} & C_{13} & 0 & 0 & 0 \\
C_{13} & C_{13} & C_{33} & 0 & 0 & 0 \\
0 & 0 & 0 & C_{44} & 0 & 0 \\
0 & 0 & 0 & 0 & C_{44} & 0 \\
0 & 0 & 0 & 0 & 0 & C_{66}
\end{array}\right|
$$

Matrix $D^{\varphi=90^{\circ}, \theta}=D$ (omitting indices $\varphi$ and $\theta$ ) in discussed case may be presented in the following way:

$$
D=\left|\begin{array}{cccccc}
d_{11} & d_{12} & d_{13} & 0 & d_{15} & 0  \tag{4}\\
d_{21} & d_{22} & d_{23} & 0 & d_{25} & 0 \\
d_{31} & d_{32} & d_{33} & 0 & d_{35} & 0 \\
0 & 0 & 0 & d_{44} & 0 & d_{46} \\
d_{51} & d_{52} & d_{53} & 0 & d_{55} & 0 \\
0 & 0 & 0 & d_{64} & 0 & d_{66}
\end{array}\right|
$$

and the elements can be expressed by means of components of tensor $C_{i j k l}$ (in Voigt's notation) and the tilt angle $\theta$ as follows:

$$
\begin{align*}
d_{11}= & C_{11} \cos ^{4} \theta+2 C_{13} \cos ^{2} \theta \sin ^{2} \theta+C_{33} \sin ^{4} \theta+ \\
& +4 C_{44}(\sin \theta \cos \theta)^{2}  \tag{5}\\
& d_{21}=d_{12}=C_{12} \cos ^{2} \theta+C_{13} \sin ^{2} \theta  \tag{6}\\
d_{31}= & d_{13}=\left(C_{11} \cos ^{2} \theta+C_{13} \sin ^{2} \theta\right) \sin ^{2} \theta+\left(C_{13} \cos ^{2} \theta+\right. \\
& \left.+C_{33} \sin ^{2} \theta\right) \cos ^{2} \theta-4 C_{44} \sin ^{2} \theta \cos ^{2} \theta  \tag{7}\\
d_{51}= & d_{15}=\left[C_{13} \cos ^{2} \theta+\mathrm{C}_{33} \sin ^{2} \theta-\mathrm{C}_{11} \cos ^{2} \theta-\right. \\
& -\mathrm{C}_{13} \sin ^{2} \theta+2 \mathrm{C}_{44}\left(\cos ^{2} \theta-\sin ^{2} \theta-\right. \\
& \left.\left.-\sin ^{2} \theta\right)\right] \sin \theta \cos \theta  \tag{8}\\
& d_{22}=C_{11} \tag{9}
\end{align*}
$$

$$
\begin{align*}
& d_{32}=d_{23}=C_{12} \sin ^{2} \theta+C_{13} \cos ^{2} \theta  \tag{10}\\
& d_{25}=d_{52}=\left(C_{13}-\mathrm{C}_{12}\right) \sin \theta \cos \theta \tag{11}
\end{align*}
$$

$$
\begin{align*}
d_{33}= & \left(C_{11} \sin ^{2} \theta+C_{13} \cos ^{2} \theta\right) \sin ^{2} \theta+\left(C_{13} \sin ^{2} \theta+\right. \\
& \left.+C_{33} \cos ^{2} \theta\right) \cos ^{2} \theta+4 C_{44} \sin ^{2} \theta \cos ^{2} \theta \tag{12}
\end{align*}
$$

$$
\begin{aligned}
d_{35} & =d_{53}=\left[\mathrm{C}_{13} \sin ^{2} \theta+\mathrm{C}_{33} \cos ^{2} \theta-\mathrm{C}_{11} \sin ^{2} \theta-\right. \\
& \left.-\mathrm{C}_{13} \cos ^{2} \theta-2 \mathrm{C}_{44}\left(\cos ^{2} \theta-\sin ^{2} \theta\right)\right] \sin \theta \cos \theta
\end{aligned}
$$

$$
\begin{equation*}
d_{44}=C_{44} \cos ^{2} \theta+C_{66} \sin ^{2} \theta \tag{14}
\end{equation*}
$$

$$
\begin{equation*}
d_{46}=d_{64}=\left(C_{44}-\mathrm{C}_{66}\right) \sin \theta \cos \theta \tag{15}
\end{equation*}
$$

$d_{55}=\left(C_{11}-2 \mathrm{C}_{13}+\mathrm{C}_{33}\right) \sin ^{2} \theta \cos ^{2} \theta+C_{44}\left(\cos ^{2} \theta-\sin ^{2} \theta\right)^{2}($

$$
\begin{equation*}
d_{66}=C_{44} \sin ^{2} \theta+C_{66} \cos ^{2} \theta \tag{17}
\end{equation*}
$$

and other remaining components are equal zero.
For small tilt angles, when $\theta \rightarrow{ }^{\circ} 0$ we have:

$$
\begin{align*}
& d_{11} \rightarrow C_{11} \\
& d_{12}=d_{21} \rightarrow C_{12} \\
& d_{13}=d_{31} \rightarrow C_{13} \\
& d_{51}=d_{15} \rightarrow 0 \\
& d_{23}=d_{32} \rightarrow C_{13} \\
& d_{25}=d_{52} \rightarrow 0 \\
& d_{33} \rightarrow C_{33}  \tag{18}\\
& d_{35}=d_{53} \rightarrow 0 \\
& d_{44} \rightarrow C_{44} \\
& d_{46}=d_{64} \rightarrow 0 \\
& d_{55} \rightarrow C_{44} \\
& d_{66} \rightarrow C_{66}
\end{align*}
$$

So for small tilt angles $\theta \rightarrow 0$, as expected, matrix

$$
\begin{equation*}
\underset{\theta \rightarrow 0}{D} \rightarrow C \tag{19}
\end{equation*}
$$

The initial point in the discussion on construction of algorithmic solutions will be Hook`s law - basic relation between the tensor of stress $T_{i j}$ and strain $E_{i j}$, which results in a conclusion that each stress component is a linear function of strain, i.e.:

$$
\begin{equation*}
T_{i j}=d_{i j k l} E_{k l}=d_{i j k l} E_{l k} \tag{20}
\end{equation*}
$$

While the tensor of strain is

$$
\begin{equation*}
E_{l k}=\frac{1}{2}\left(U_{l, k}+U_{k, l}\right) \tag{21}
\end{equation*}
$$

Presenting the relation (20) in detail, we receive:

$$
\begin{align*}
T_{i j}= & d_{i j 11} E_{11}+d_{i j 22} E_{22}+d_{i j 33} E_{33}+2 d_{i j 23} E_{23}+ \\
& +2 d_{i j 13} E_{13}+2 d_{i j 12} E_{12} \tag{22}
\end{align*}
$$

Substituting $i, j=1,2,3$, we obtain all the components of stress tensor: $T_{11}, T_{22}, T_{33}, T_{23}=T_{32}, T_{13}=T_{31}, T_{12}=T_{21}$. In the matrix notation this is as follows:

$$
\left|\begin{array}{c}
T_{11}  \tag{23}\\
T_{22} \\
T_{33} \\
T_{32}=T_{23} \\
T_{31}=T_{13} \\
T_{12}=T_{21}
\end{array}\right|=D\left|\begin{array}{c}
E_{11}=E_{x x}=U_{x, x} \\
E_{22}=E_{y y}=U_{y, y}=0 \\
E_{33}=E_{z z}=U_{z, z} \\
2 E_{23}=2 E_{y z}=U_{y, z}=0 \\
2 E_{31}=U_{z, x}+U_{x, z} \\
2 E_{12}=U_{x, y}=0
\end{array}\right|
$$

In relation (23) it has been considered that derivatives of the field with regard to coordinate $y$ equal zero. Starting with the general law of movement (disregarding the external force)

$$
\begin{equation*}
T_{i j, j}=\rho \frac{\partial^{2} U_{i}}{\partial t^{2}} \tag{24}
\end{equation*}
$$

where $\rho$ is the medium density, and $t$ denotes time, let us write equations for the horizontal $U_{x}\left(U_{1}\right)$ and vertical $U_{z}\left(U_{3}\right)$ component

$$
\begin{align*}
& T_{11,1}+T_{13,3}=\rho \frac{\partial^{2} U_{1}}{\partial t^{2}}  \tag{25a}\\
& T_{31,1}+T_{33,3}=\rho \frac{\partial^{2} U_{3}}{\partial t^{2}} \tag{25b}
\end{align*}
$$

Using the matrix equation (23) with relation to equations (25), the following relations are received:

$$
\begin{align*}
d_{11} U_{x, x x} & +d_{55} U_{x, z z}+2 d_{15} U_{x, x z}+d_{15} U_{z, x x}+ \\
& +\left(d_{13}+d_{55}\right) U_{z, z x}+d_{53} U_{z, z z}=\rho \frac{\partial^{2} U_{x}}{\partial t^{2}}  \tag{26a}\\
d_{51} U_{x, x x} & +\left(d_{31}+d_{55}\right) U_{x, x z}+d_{35} U_{x, z z}+d_{33} U_{z, z z}+ \\
& +2 d_{35} U_{z, x z}+d_{55} U_{z, x x}=\rho \frac{\partial^{2} U_{z}}{\partial t^{2}} \tag{26b}
\end{align*}
$$

Adopting for small tilt angles $\theta d_{15}=d_{51} \approx 0$ and $d_{53}=d_{35} \approx 0$ and applying Fourier`s transformation \(\left(x \rightarrow k_{x} z \rightarrow k_{z} t \rightarrow \omega\right) z\) for equations (26) we obtain matrix equation analogical to Christoffel`s relation:

$$
\left(\begin{array}{ll}
d_{11} k_{x}^{2}+d_{55} k_{z}^{2}-\rho \omega^{2} & \left(d_{13}+d_{55}\right) k_{x} k_{z}  \tag{27}\\
\left(d_{31}+d_{55}\right) k_{x} k_{z} & d_{33} k_{z}^{2}+d_{55} k_{x}^{2}-\rho \omega^{2}
\end{array}\right)\left|\begin{array}{l}
U_{x} \\
U_{z}
\end{array}\right|=0
$$

In equation (27), $k_{x}$ and $k_{z}$ denote wavenumbers in horizontal and vertical direction, whereas $\omega$ is frequency.

It should be noted that matrix equation (27) for medium VTI, i.e. in the case when the tilt angle $\theta=0$ transforms into analogical equation derived by Q. Han and R.S. Wu [5]. Equation (27) results in dispersion relation

$$
\begin{equation*}
b_{0} k_{z}^{4}+b_{1} k_{z}^{2}+b_{2}=0 \tag{28}
\end{equation*}
$$

where:

$$
\begin{align*}
& b_{0}=d_{33} d_{55} \\
& b_{1}=k_{x}^{2}\left(d_{11} d_{33}-d_{13}^{2}-2 d_{11} d_{55}\right)-\rho \omega^{2}\left(d_{33}+d_{55}\right) \\
& b_{2}=k_{x}^{4} d_{11} d_{55}-\rho \omega^{2}\left(d_{11}+d_{55}\right) k_{x}^{2}+\rho^{2} \omega^{4} \tag{29}
\end{align*}
$$

In general, equation (28) has four solutions corresponding to longitudinal wave qP and shear wave qSV (polarized in plane $x-z$ ) forward and backward propagation. Q. Han and R.S. Wu [5] concluded on the basis of numerous experiments that wave velocity qSV does not provide significant contribution in the quantity of vertical wavenumber of longitudinal wave qP , therefore it can be assumed that velocity of wave qSV equals zero [1]. On assumption that $b_{0}=0$, we receive:

$$
\begin{align*}
& b_{1}^{\prime}=k_{x}^{2}\left(d_{11} d_{33}-d_{13}^{2}\right)-\rho \omega^{2} d_{33} \\
& b_{2}^{\prime}=\rho^{2} \omega^{4}-\rho \omega^{2} d_{11} k_{x}^{2} \tag{30}
\end{align*}
$$

Let expressions $b_{1}^{\prime}$ and $b_{2}^{\prime}$ be represented by parameters analogical to Thomsen`s, i.e.:

$$
\begin{gather*}
\varepsilon=\frac{d_{11}-d_{33}}{2 d_{33}}  \tag{31}\\
\partial=\frac{\left(d_{13}+d_{55}\right)^{2}-\left(d_{33}-d_{55}\right)^{2}}{2 d_{33}\left(d_{33}-d_{55}\right)} \tag{32}
\end{gather*}
$$

hence

$$
q=1+2 \varepsilon=\frac{d_{11}}{d_{33}}
$$

and for $d_{55}=0$

$$
1+2 \partial=\frac{d_{13}^{2}}{d_{33}^{2}}
$$

Assuming that the velocity of longitudinal wave $V_{p}=\left(\frac{C_{33}}{\rho}\right)^{1 / 2}$ a $d_{33}=C_{33} \cos ^{2} \theta$ we receive:

$$
\begin{equation*}
d_{33}=\rho V_{p}^{2} \cos ^{2} \theta=\rho V_{p p}^{2} \tag{33}
\end{equation*}
$$

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Similarly, using parameters $\varepsilon$ and $q=1+2 \varepsilon$ and $\eta=2(\varepsilon-\delta)$ we have:

$$
\begin{equation*}
d_{11} d_{33}=(1+2 \varepsilon) d_{33}^{2}=q \rho^{2} V_{p}^{4} \cos ^{4} \theta=q \rho^{2} V_{p p}^{4} \tag{34}
\end{equation*}
$$

where: $V_{p p}=V_{p} \cos \theta$ denotes the velocity of longitudinal wave along the vertical axis

$$
\begin{gather*}
d_{13}^{2}=(1+2 \partial) d_{33}^{2}=(1+2 \partial) \rho^{2} V_{p p}^{4}  \tag{35}\\
d_{11}=(1+2 \varepsilon) \rho V_{p p}^{2} \tag{36}
\end{gather*}
$$

From relation (29) and relations (30)-(35) we obtain the expression for vertical component of wavenumber $k_{z}$

$$
\begin{equation*}
k_{z}=\left(\frac{S_{p}^{4} \omega^{4}-q S_{p}^{2} \omega^{2} k_{x}^{2}}{S_{p}^{2} \omega^{2}-\eta k_{x}^{2}}\right)^{1 / 2} \tag{37}
\end{equation*}
$$

where: $S_{p}=\frac{1}{V_{p p}}$ denotes vertical slowness.
It is not difficult to notice that parameters $q$ and $\eta$ which can be seen in expression (37) are represented by their values in model VTI, therefore for horizontally laminated medium

$$
\begin{align*}
& q=q_{V T I} \cdot \cos ^{2} \theta \\
& \eta=\eta_{V I I} \cdot \cos ^{2} \theta \tag{38}
\end{align*}
$$

where: $q_{V T I}$ and $\eta_{V T I}$ denote the values of these parameters in VTI medium.

Thus relation (37) for medium TTI is analogical to the relation derived by Q. Han and R.S. Wu [5] for the horizontally laminated medium VTI. In order to use it in the process of propagation and migration of compressional waves one should know parameters $q$ and $\eta$, the velocity of P waves along the axis and the tilt angle $-\theta$, the plane of isotropy. We will apply received wavenumber $k_{z}$ for
migration $\mathrm{MG}(\mathrm{F}-\mathrm{K})$ in the domain of wavenumbers and frequencies and spatial coordinates $x$ and time $t[8,11]$. Migration process performed in this way occurs in two stages. At the first stage, the relocation of the wave field takes place

$$
\begin{equation*}
U^{\prime}\left(k_{x}, z_{j}+\Delta z, \omega\right)=e^{-i k z_{0} \Delta z} U\left(k_{x}, z_{j}, \omega\right) \tag{39}
\end{equation*}
$$

from the level of $z_{j}$ to $z_{j}+\Delta z$ by means of exponential operator with vertical wavenumber $k z_{0}$ corresponding to a homogeneous medium. At the second stage, correction of the wave field follows $U^{\prime}\left(x, z_{j}+\Delta z, \omega\right)$ - Fourier transforms $\left(k_{x} \rightarrow \mathrm{x}\right)$ of the field $U^{\prime}\left(k_{x} z_{j}+\Delta z, \omega\right)$ by way of spatial filter $F_{j}(x, \omega)=\left[1-i / 2 \Delta z M_{j(x)}\right]^{-1}$, which is the sum of Neumann's power series

$$
\begin{equation*}
M_{j(x)}=\Sigma k_{z_{0}}^{-1}\left(k_{z_{0}}^{2}-k_{z}^{2}\right) e^{i k_{x} x} d k_{x} \tag{40}
\end{equation*}
$$

This relation may be represented in this way:

$$
\begin{equation*}
U_{z j+\Delta z}+=F_{j}(x, \omega) U^{\prime}\left(x, z_{j}+\Delta z, \omega\right) \tag{41}
\end{equation*}
$$

The correction positions the wave field in the function of spatial coordinates, taking into account the differences between parameters of a homogenous medium and parameters of heterogeneous medium in the function of lateral coordinates. For the prestack migration, the algorithm of extrapolation will be the product of corrective functions $F$ related to the sources and receivers, while corrected field $U^{\prime}$ will be a function of coordinates of sources and receivers. With the zero-offset migration in relation (37) slowness $S_{p}$ must be multiplied by 2 . Further steps follow in an analogical way as in model VTI and it was discussed in detail in the article by A. Kostecki [11]. It should be noted that even when the elastic parameters are independent from the spatial coordinates, it is essential to take into consideration the differences in vertical wavenumbers as a result of different tilt angles of the laminated medium.

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## Literature

[1] Alkhalifah T.: Acoustic approximation for processing in transversely isotropic media. Geophysics, 63, 623-631, 1998.
[2] Auld B.A.: Acoustic fields and waves in solids. Krieger Publishing Company, vol. 1, 2, 1990.
[3] Bansal R., Sen M.: Finite - difference modelling of S-wave splitting in anisotropic media. Geophysical Prospecting 56, 293-312, 2008.
[4] Crampin S., Chesnokov E, Hipkin R.: Seismic anisotropy - the state of the art. Geophys. J. Roy. Astr. Soc., 76, 1-16, 1984.
[5] Han Q., Wu R.S.: A one-way dual domain propagator for scalar qP waves in VTI medium. Geophysics, vol. 70, D9-D-17, 2005.
[6] Issack I., Lavton D.: A practical method for estimating effective parameters of anisotropy from reflection seismic data. Geophysics, vol. 69, 681-689, 2004.
[7] Jianlin Zhu, Ji Dorman: Two-dimensional, three-component wave propagation in a transversely isotropic medium with arbitrary orientation - finite - element modeling. Geophysics, vol. 65, no 3, pp. 934-942, 2000.
[8] Kostecki A., Półchłopek A.: Seismic migration before sum-

## artykuły

ming up the wavefield in a medium with lateral velocity heterogeneities. Studies by Oil and Gas Institute, no 94, 1998.
[9] Kostecki A., Półchłopek A.: Stable depth extrapolation of seismic wavefields by Neumann series. Geophysics, 63, 2063-2071, 1998.
[10] Kostecki A., Półchłopek A.: A study of structural reproduction in anisotropic medium VTI. Prace INiG, no 137, pp. 39-60, 2006.
[11] Kostecki A.: Algorithms of depth migration in anisotropic medium VTI. Nafta-Gaz, no 11, 661-666, 2007.
[12] Postma G.M.: Wave propagation in a stratified media. Geophysics, vol. 20, no 4, 1955.
[13] Schoenberg M.A.: Seismic characterization of reservoirs containing multiple fracture sets. Geophysical Prospecting, vol. 57, no 2, pp. 169-186, 2009.
[14] Thomsen L.: Weak elastic anisotropy. Geophysics, vol. 51, 1954-1966, 1986.


Andrzej KOSTECKI - Professor of geophysics. The main subject of interest - electromagnetic and seismic wave propagation, reproduction of deep geological structures by means of seismic migration, the analysis of migration velocities, seismic anisotropy. The author of 130 publications.

## ZAKŁAD TECHNIKI STRZELNICZEJ INiG W KRAKOWIE

OFERTA NA WYKONYWANIE BADAŃ ODPORNOŚCI CIŚNIENIOWEJ I TERMICZNEJ URZADZEŃ
Oferujemy wykonywanie badań odporności ciśnieniowej urządzeń w warunkach podwyższonej temperatury i w temperaturze otoczenia na stanowisku termobarycznym INiG w Krakowie.

- maksymalne wymiary gabarytowe:

| - długość | 1850 mm |
| :--- | :--- |
| - średnica | 140 mm |
|  | 100 kG <br>  <br>  <br>  <br>  <br>  <br>  <br>  <br>  <br>  <br>  <br>  <br>  <br>  <br>  <br>  <br>  |

- maksymalny ciężar:
- ciecz robocza:
- maksymalne ciśnienie badania:
- maksymalna temperatura badania:
$180^{\circ} \mathrm{C}$
- ogrzewanie pośrednie w płaszczu olejowym,
- rejestracja ciagła ciśnienia i temperatury,
- możliwe wykonywanie badań „do zniszczenia" w przypadku urządzeń o niewielkiej objętości,
- możliwość wyprowadzenia sygnału elektrycznego z badanego urządzenia linią 2-przewodową.

Kierownik: mgr inż. Antoni Frodyma
Adres: ul. Bagrowa 1, 30-733 Kraków
Telefon: 12 653-25-12 wew. 137
Faks: 12 653-16-65
E-mail: antoni.frodyma@inig.pl


