

Use of MLCM3 Software for Flash Flood Modeling and Forecasting

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ABSTRACT

Accurate and timely flash floods forecasting, especially, in ungauged and poorly gauged basins, is one of the most important and challenging problems to be solved by the international hydrological community. In changing climate and variable anthropogenic impact on river basins, as well as due to low density of surface hydrometeorological network, flash flood forecasting based on “traditional” physically based, or conceptual, or statistical hydrological models often becomes inefficient. Unfortunately, most of river basins in Russia are poorly gauged or ungauged; besides, lack of hydrogeological data is quite typical. However, the developing economy and population safety necessitate issuing warnings based on reliable forecasts. For this purpose, a new hydrological model, MLCM3 (Multi-Layer Conceptual Model, 3rd generation) has been developed in the Russian State Hydrometeorological University. The model showed good results in more than 50 tested basins.

Keywords: hydrological model, floods forecasting, river basins, changing climate.

INTRODUCTION

In the Russian Federation, approximately 40 to 68 major floods occur each year. The largest floods occurred in 2012 in Krasnodar Territory and in the Republic of Adygea in 2017.

The need to protect the rapidly developing Russia's economy and population requires timely warnings based on reliable forecasts. While such warnings should be issued for about one of approximately 2.6 million Russian streams (most of which are ungauged or poorly gauged), this can be done only in fully automated and quite approximate (virtually, qualitative) mode. In changing climate and variable anthropogenic impact on river basins, as well as due to quite low density of the surface hydrometeorological network, flash flood forecasting based on “traditional” physical-

ly based, or conceptual, or statistical hydrological models often becomes inefficient [Pivovarova 2016]. Thus, the basic task of the presented study can be posed as developing a fully automated system of qualitative forecasting risks of flooding regardless of the quality and spatiotemporal resolution of the available hydrometeorological data.

The authors of the presented study realize that accurate and timely flash floods forecasting, especially, in ungauged and poorly gauged basins, is one of the most important and challenging problems to be solved by the international hydrological community; also, they understand that the accuracy and efficiency of forecasting depend, first of all, on 1 – data resolution and quality and 2 – the ability of used models to reflect the runoff generation processes properly. Therefore, the efficiency of flood forecasting in different basins also

will be different, but, while we are keen on predicting the overall risk of flooding (say, in terms of its probability and level of danger), we agree, by default, that the streamflow forecasting carried out for a big number of streams cannot be as accurate as that for a single stream, and should suppose presenting forecasts in the same format (for instance, the flood risk level).

The hydrological models that would be suitable for implementation in automated mode should be:

- easily, fast and efficiently calibrated (and recalibrated, if necessary);
- inexpensive in terms of computational/processor resources [Phong et al. 2015];
- forced by commonly available data (such as precipitation and evaporation/evapotranspiration), and
- reliable enough even when being forced by “bad” data.

For this purpose, a new hydrological model, MLCM3 (Multi-Layer Conceptual Model, 3rd generation) has been developed in the Russian State Hydrometeorological University.

It should be noted that a large part of hydrological studies, monographs and articles are devoted to hydrological modeling, as well as the development of new hydrological models and the improvement of the existing ones.

The article [Kuzmin 2009] initiates a series of three articles on the automatic calibration of operational multi-parameter models used for flash flood forecasting in an automated regime. The first article of the series contains a critical analysis of the “mainstream“ in hydrological models calibration and presents the basic principles of the Stepwise Line Search (SLS) algorithm and its modifications, practically feasible and robust parametrization approaches that would be suitable for an automated systems used for flash flood forecasting.

In the work [Alfonso 2010] written by leading European experts in the field of flood and flash flood modeling, the latest studies of the effectiveness of optimizations of various versions of lumped, distributed and half-distributed hydrological models, are summarized. In the Russian Federation, conclusions made in this paper can be used as a methodological basis for the development of models in basins with a low spatial-temporal discreteness of ground-based observations, so individual ideas of the article were

used in developing the calibration method for the MLCM3 model.

It is known that the models used by the forecasting organizations should be both reliable and simple, but also be able to provide sufficient lead-time for warnings and the desired degree of accuracy. In general, the choice of the model depends on the following factors: the amount of data available, the complexity of the hydrological processes to be modeled, the reliability required, the accuracy and lead-time, the type and frequency of flooding, and user requirements [Refsgaard 1997].

In many cases, for hydrological modeling and forecasting, when projections of expenditure or water levels along the river are needed, hydrometeorology specialists use the models of “rainfall-runoff” type in conjunction with the models of the propagation of the flood wave. If the precipitation is presented in the form of snow, snowmelt models are used. These models differ in terms of accuracy and complexity, occupying the entire spectrum from the models based on the use of the index of previous humidification, to multiparameter conceptual models or process models [WMO. Guide to Hydrological Practices 1994].

In our case study, in which we had to find a certain compromise between various extremes, conceptual models seem to be preferable. Typically, the conceptual, hydrological models include the description of the processes of water supply, loss, redistribution of moisture in the soil, surface, soil and subterranean inflow to the channel network, as well as their transformation into the hydrograph. However, the degree of detail in the description of these processes, as well as the dependencies used in various models, can vary significantly and lead to their diversity.

This type of models is operationally implemented in many countries. For example, the Hydrologiska Byråns Vattenbalansavdelning model (HBV) was developed by Bergström at the Swedish Meteorological and Hydrological Institute for modeling and analysis of river runoff. The HBV model is a conceptual model that converts precipitation, atmospheric temperature and potential natural water loss into either snow melting, or into runoff or inflow into the reservoir. The HBV model can be considered as a model with half-distributed parameters: the catchment area is divided into private catchments, and the altitudinal zoning method is also applied [Johansson 1997].

Additionally, there is another well-known conceptual model of MGB-IPH, developed by the

research group IPH-UFRGS, which specializes in large-scale hydrology. This model is used to work with many South American basins. MGB-IPH is a sediment-drain model that uses precipitation, air temperature, relative humidity, wind speed, insolation and atmospheric pressure to calculate the flow rate in a given basin.

The well-known Sacramento Soil Moisture Accounting Model (SAC-SMA) [Burnash et al. 1973, Reed et al. 2005] is successfully implemented in the United States for flash flood forecasting for four decades. This model is very popular among hydrologists. The original version of the Sacramento model includes 16 optimized parameters, the values of which can be estimated by observing soil moisture and entering data on soil type; however, in most cases, it is possible to replace 5–7 parameters with constants without any modeling quality loss.

FEATURES OF THE CONCEPTUAL HYDROLOGICAL MODEL MLCM3

In this paper, a basic version of the MLCM3 is presented. MLCM3 is a rainfall-runoff model with a flexible structure and high level of “conceptualization”. The model forcing includes precipitation and evaporation data, basically coming from NWP model output. Water comes to the outlet through several layers; their number, as well as two parameters (thickness and infiltration rate) for each of them, surface flow velocity (when the top layer is full of water) are optimized.

MLCM3 has a variable number of parameters: a number of layers N , depth of i^{th} layer Z_p , infiltration rates α_i (which can be constant or computed using individual equations for each layer), and velocity of the surface runoff α_0 . Thus, a total number of parameters will be equal to $2N + 2$. If the unit hydrograph is approximated with 2-parameter Gamma-distribution, two extra parameters are added. Furthermore, Muskingum-Cunge, or kinematic wave, or any other routing scheme can be applied; in our case, Muskingum-Cunge algorithm is used by default. The Muskingum routing procedure relies on the use of the storage-outflow relationship $Q_2 = C_1 I_1 + C_2 I_2 + C_3 Q_1$, where Q_j is the outflow rate; I_k is the inflow rates coming from upstream; coefficients C_1 , C_2 and C_3 depend on the Muskingum parameters X and K that may be optimized or treated as empirical and fitted when an outflow hydrograph is available at

a gauge. A priori estimates of X and K can be also obtained by using the physical channel properties x , c , b and s . Figure 1 shows the scheme of the conceptual hydrological model MLCM3 and the Table 1 present the forcing data for it [Barrett 2008, Kuzmin et al. 2008].

All the parameters can be optimized; there are many possible strategies to use the available processor resources for this purpose. For instance, for urban catchments, a number of layers can be fixed (up to 2), and all possible parameter sets can be examined. If some moisture holding data are available, a certain number of parameters can be also fixed, while the rest of them will be optimized. Alternatively, soil moisture content of the top layer can be determined by using the satellite observations (e.g. soil moisture data from the AMSR-E), while the rest of layers will be parameterized separately, etc.

Time step may also vary. There are two default options (1 h and 24 h), but a modeler can set any value depending on his/her needs.

INTER-COMPARISON OF THE MLCM3 MODEL WITH THE SACRAMENTO MODEL

The conceptual hydrological model of the Sacramento Soil Moisture Accounting Model (SAC-SMA) is one of the most well-known and most commonly used hydrological models in modeling and forecasting. Only the standard hydrometeorological network observations, climatological information and publicly available cartographic data are used as the “input”. However, such a hydrological model is extremely difficult to calibrate if there is insufficient measurement data (shortage of posts and lack of measurement of some quantities are a fact in the Russian Federation).

Inter-comparison of MLCM3 and SAC-SMA demonstrated their similar efficiency in the flash flood modeling based on the radar precipitation data. However, unlike SAC-SMA, MLCM3 was calibrated by using SCE algorithm and implemented without soil data. The obtained results for test catchments are presented in Table 2.

As can be seen from Figure 1 and the results of numerical experiments in Table 2, MLCM3 is a bit simpler than the Sacramento model. However, it is more flexible than SAC-SMA, because its concept supposes propagation of water in N layers, while in SAC-SMA this process is limited

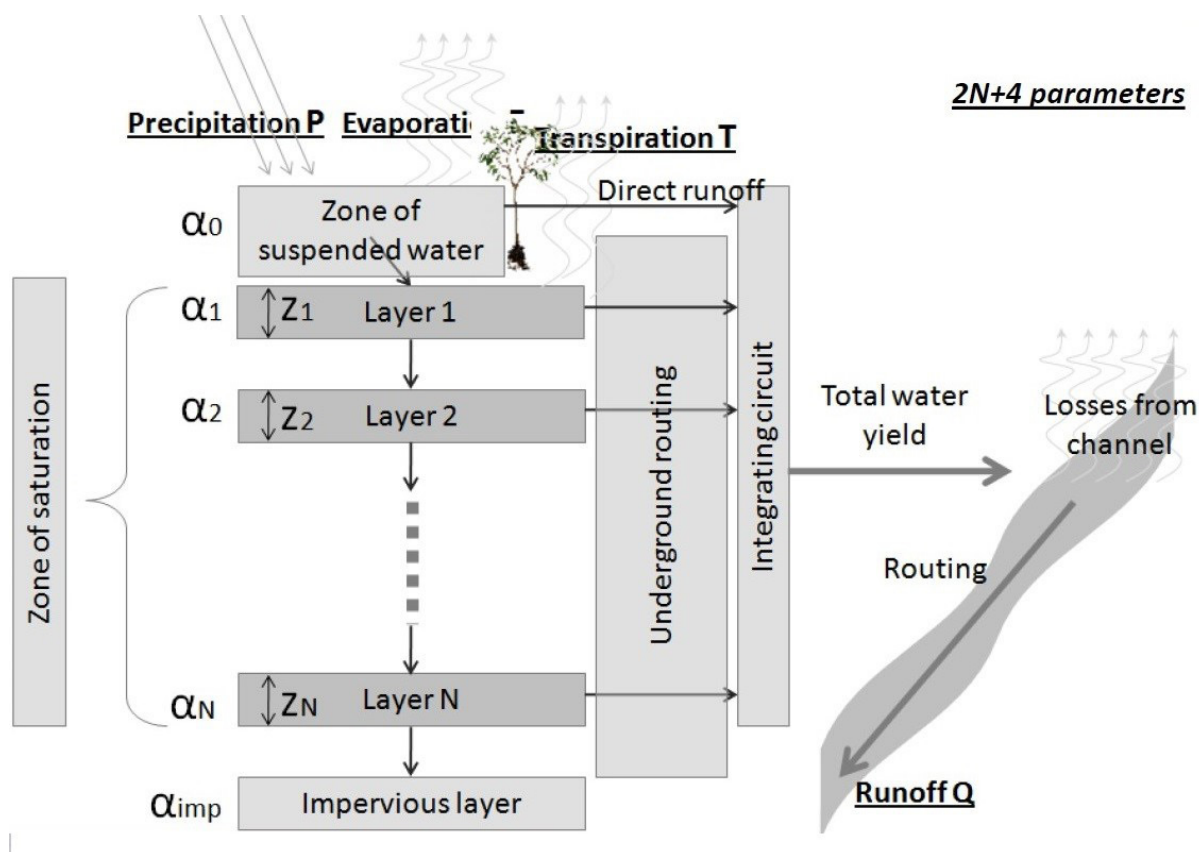


Figure 1. The conceptual hydrological model MLCM3

by 5 storages and MLCM has a higher level of conceptualization, than SAC-SMA. Unlike SAC-SMA, it does not require any soil data (however, it can be used if such data are available).

The main advantage of the MLCM3 model is its ability to use it with a smaller amount of input information than is required in the case of the Sacramento model. The main features of the MLCM3 model, in contrast to the Sacramento model, are its high predictive efficiency and the performance of calibration and validation procedures in a fully automated mode.

MLCM3 SOFTWARE FOR FLASH FLOOD FORECASTING

The software «MLCM3» is intended for forecasting and modeling rainfall flash floods. The MLCM3 is an application that utilizes an intuitive user interface that makes imputing and editing experiments with data fast and efficient. This program uses a client/server based model. The client is what the user uses to add researcher's records as well as edit them. The client program will communicate with a university server

Table 1 Forcing data for MLCM3

Data type	Symbol
Precipitation	P
Evapotranspiration	ET
Catchment area	S
Main channel length	L
Actual runoff (for calibration only)	Q

(RSHU) that saves all the information for each researcher (information about experiment results) and includes some of their personal information (login and password).

The main function of the MLCM3 software is the processing of input files containing data on hydrometeorological variables at the time of the calculation. As a background forecast, the data of the external mesoscale meteorological model (Weather Research and Forecasting (WRF) – a package of modeling meteorological processes) are used. The program is designed for forecasting the river flow, which allows assessing the hydrometeorological risks in the conditions of climate change and variable anthropogenic load on river basins.

Table 2. The results of numerical experiments on modeling of rainfall flash floods

No.	Basin	Watercourse code	Value of the objective function MSOF	
			MLCM3	SAC-SMA
1	Austin – Onion Creek	08159000	18.87	19.36
2	Justin – Denton Creek	08053500	14.72	16.13
3	Houston – Greens Bayou	08076000	11.03	11.35
4	Georgetown – South Fork San Gabriel	08104900	14.80	16.22
5	Greenville – Cowleech Creek	08017200	13.70	14.39
6	Houston – Brays Bayou	08075000	25.65	27.02
7	Hunt – Guadalupe River	08165500	28.91	30.99
8	Justiceburg – Double Mt Fork	08079600	11.95	12.19

There are two different types of users: Admin Users and Users. Admin Users are MLCM3 software developers who have the authority to give Users the access to use the software. On the other hand, users have the right to work with the software and upload information to it. In order to get started with the software, one needs to go to the site <http://mlcm2.rshu.ru/> and log in to their account with the login and password received in advance from the administrator. The Figure 2 shows the interface of the MLCM3 software.

The top tabs are optional and include:

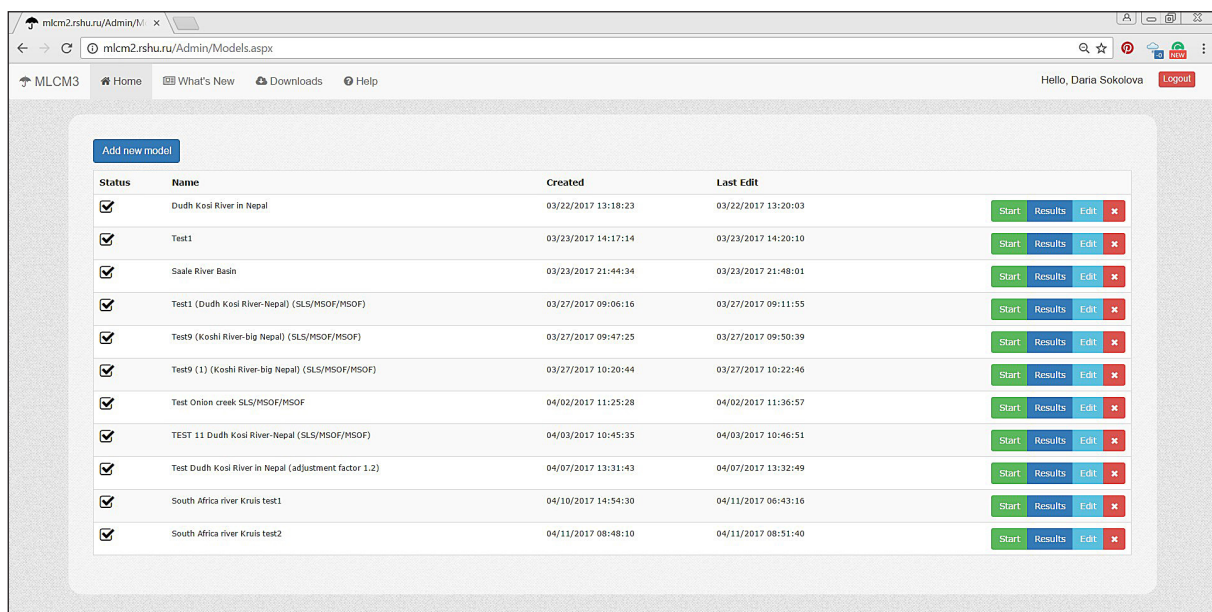
- In the tab “**Home**” – store experiments.
- In the tab “**What’s New**” – read information about the program’s improvements.
- In the tab “**Downloads**” – download a converter for data sets.
- In the tab “**Help**” – download User’s Manual.

In MLCM3, it is possible to carry out a full technical adjustment of the model parameters.

One can choose from three types of calibration and pre-calibration. The user can perform the SLS calibration (Stepwise Line Search), the Nelder-Mead methods and the extended Nelder-Mead. Furthermore, the user can select the type of the objective function and the validation. The choice includes: the root-mean-square error, the criterion S/σ and the multi-scale objective function MSOF. In order to emulate this multi-scale nature of manual calibration, we employ an objective function composed of contributions from a wide range of time scales of aggregation. The particular objective function MSOF used in this work was proposed by Koren and has the following form:

$$J = \sqrt{\sum_{k=1}^n \left(\frac{\sigma_1}{\sigma_k} \right)^2 \sum_{i=1}^{m_k} (q_{o,k,i} - q_{s,k,i}(X))^2}$$

where: $q_{o,k,i}$ and $q_{s,k,i}$ – measured and modeled discharge of water averaged over the time interval k

**Figure 2.** The interface of the MLCM3 software

σ_k – Mean square error deviation of the discharge of water rate of k
 n – Total number of scales,
 m_k – the number of items of each scale k
 where $q_{o,k,i}$ and $q_{s,k,i}$ k , are the observed and simulated flows averaged over time interval k ; σ_k is the standard deviation of discharge at that scale; n is the total number of scales used, and m_k is the number of ordinates at the scale k .

This weighting scheme assumes that the uncertainty in modeled streamflow at each scale is proportional to the variability of the observed flow at that scale. Another important advantage of using the multi-scale objective function (MSOF) is that it smoothes the objective function’s surface, and, hence, reduces the likelihood of the search getting stuck in tiny “pits”.

In Figure 4, tabs are provided for the mapping of a single stream. All buttons are optional and simulation, calibration and validation periods can be set manually in cells or graphically by hovering on a graph. After selecting all the necessary characteristics and loading data, the software processes the data, depending on the length of the time series, which can take from 2 to 10 minutes, on average.

All experimental results are available in the main menu of the MLCM3 software, users can work with them repeatedly and make changes.

EXPERIMENTAL STUDY OF THE APPLICABILITY OF THE NEW MLCM3 SOFTWARE

It is necessary to emphasize once again that using the same criterion for the accuracy of forecasts or the effectiveness of a predictive technique for model calibration and validation does not at all guarantee the best result from a practical point of view. Therefore, when calibrating the model, it must be remembered that the calibration method, the type of the objective function and the selected training samples should provide the most accurate reflection of the different phases of the hydrological regime (including rising levels, peak flood or flood, level drop and low runoff) with a different order of alternation.

In order to solve the problem of parametrization of models in the conducted experiments, we used the method of quasi-local optimization in the physically pre-determined region of the domain of definition of the objective function – SLS (Step-wiseLineSearch). The basic algorithm has several

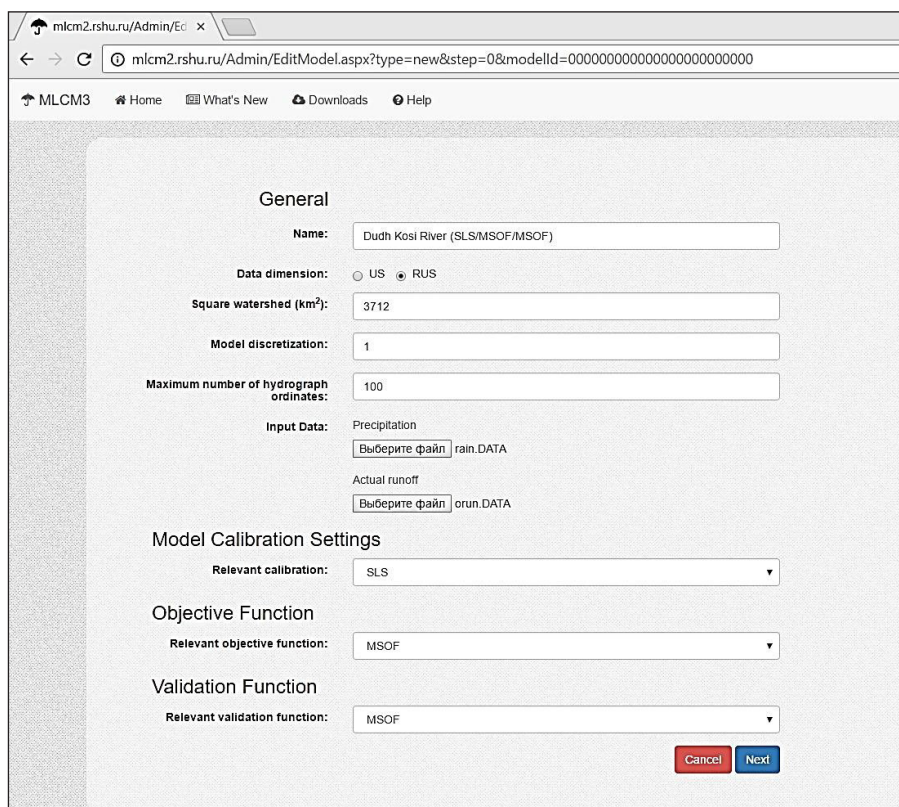


Figure 3. The software tab intended for the adjustment of technical parameters of model calibration

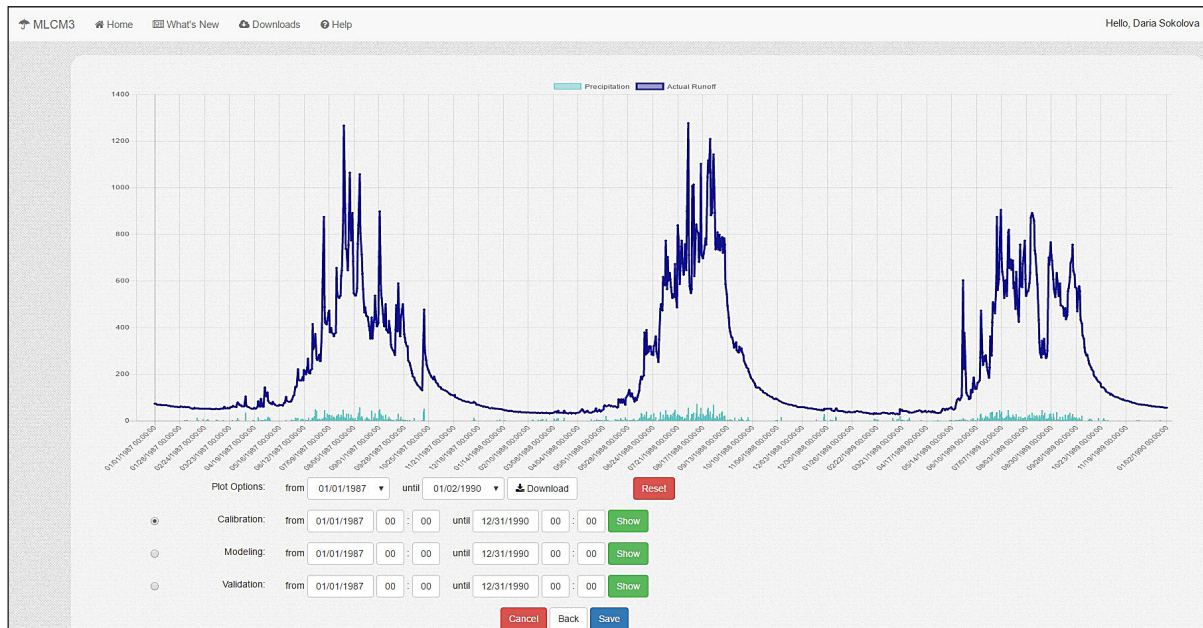


Figure 4. The software tab designed to display the flow of a single stream

modifications designed to calibrate the prognostic models used to predict runoff under different conditions. In addition, it is the functional basis for postprocessing forecasts; thus, the SLS method is a convenient integrated tool used to account for all kinds of uncertainty (“noise”) that affects the simulation results. The SLS method is most suitable in the cases when the a priori (initial) point for quasilocal optimization is correctly set and the catchment area is sufficiently well illuminated by hydrometeorological observations [Kayastha et al. 2013, Kuzmin 2001a, 2001b].

A series of experiments were conducted in the basins of Kosi region (Nepal) for 2 years in the period from 1997 to 1998. Experiment 1 was produced according to the catchment area of the Sagarmatha Post, the Dudh Kosi River and was conducted without changes in the initial parameters in the software. For the target function, a multi-scale MSOF target function was adopted. The validation by the MSOF target function is 13.2, and the model call is 5829. The SLS calibration was used. As can be seen in the graph (Fig.5), the flood recession is modeled well enough. This is relevant since it is the most accurate reflection of the different phases of the hydrological regime that is most important.

Experiment 2 was produced according to the catchment area of Kosi region (Nepal), the Sagarmatha Post, the Dudh Kosi River and was conducted without changes in the initial parameters in the software. Criteria and validation criteria

S/σ and $S/\sigma\Delta$ were used for the objective function and equal to 0.70. Calibration was performed using the Nelder-Mid method. As can be seen in the graph (Fig. 6), the flood recession is modeled somewhat worse for the selected period.

Importantly, the use of the MSOF functions in the calibration (for example, instead of the standard deviation, criteria $S/\sigma\Delta$ and $S/\sigma\Delta$ and, as well as the Nash-Sutcliffe criterion), allows the user to pay attention to a very important effect: the simulated hydrograph approximates the actual value as much as using these standard quadratic metrics [Sokolova and Kuzmin 2017].

CONCLUSION

On the basis of the numerical experiments of the studied basins in Kosi (Nepal) based on the MLCM3 software, it can be concluded that the MSOF objective function is the best suited for both validation and calibration. It has been established that the simulated hydrograph, when calibrating the MSOF objective functions, approximates the actual hydrograph as much as possible using these standard quadratic metrics, which is explained by the nature of the MSOF manual calibration and its ability to reflect different flow rates. In addition, it is important to note that the MSOF objective function provides an additional smoothing of the response surface of the objective function, which reduces the probability of

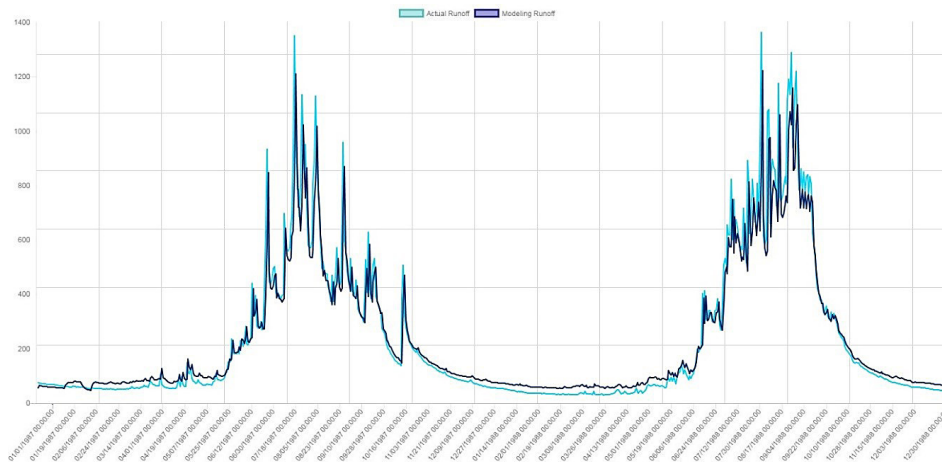


Figure 5. Dudh Kosi, Kosi region (Nepal), from 01/01/1987 to 31/12/1988; Numerical experiments: SLS calibration, MSOF objective function

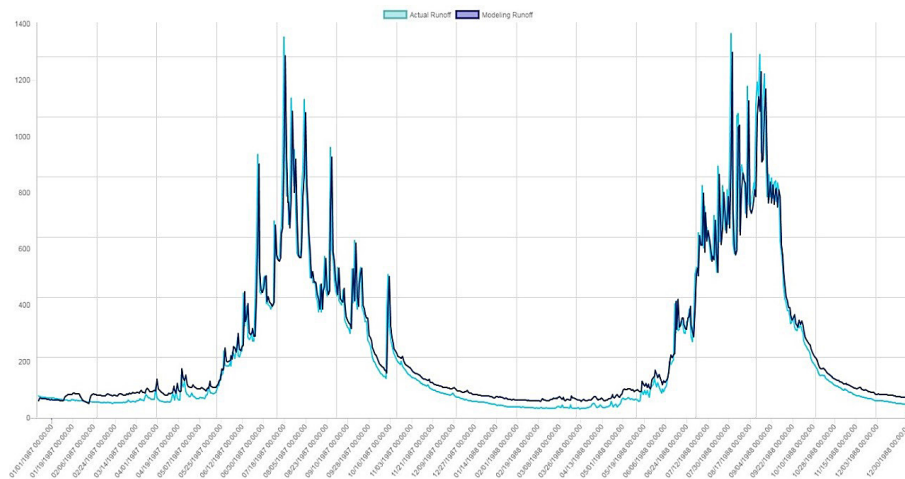


Figure 6. Dudh Kosi, Kosi region (Nepal), from 01/01/1987 to 31/12/1988; Numerical experiments: Nelder-Mead, criteria $S/\sigma\Delta$

“getting stuck” in the search for the optimum in insignificant “depressions”.

A method has been developed for presenting the results of the background forecasting of rain floods due to the fact that the applicant developed new MLCM3 software, based on the developments of the RSHU and improved in the project “Development of methodological bases and management technologies Water resources of river systems insufficiently illuminated by hydrometeorological observations (on the example of the Mekong river basin)».

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