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BULK GEOCHEMICAL DATA OF FOSSIL WOOD FROM THE MIDDLE JURASSIC CLAYS OF POLAND

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Abstract. Macroscopic observations, microscopic studies and literature data revealed that Middle Jurassic wood from Częstochowa area has a different state of preservation, and various types and degrees of mineralization and oxidation. Obtained results of organic matter fractionation illustrate a clear domination of polar fraction in the obtained extracts revealing low thermal maturity stage. Total organic carbon (TOC) values for analysed samples are in a wide range from 1.06% to 68.50%. The highest amount of TOC were measured in not or poorly mineralized wood samples but most of them are mineralized wood fragments, showing the TOC values in the range of 2% - 10%. Percentage content of carbonate in fossil wood constitute in a wide range from less than 1% CaCO₃ to above 85% CaCO₃. The resulting percentage of the total sulfur content is very varied and do not show convergence with other data such as TOC, carbonate content, etc and is most probable connected with pyritisation range. Unlike the Middle Jurassic clay samples, where long-chain and short-chain *n*-alkanes occur in similar concentrations, in wood samples always short-chain *n*-alkanes dominated, in the range from 15 to 23 carbon atoms in molecule. The values of the CPI are generally higher than 1 which indicates the contribution of organic matter derived from higher plant waxes, which are characteristic of e.g. needles from gymnosperm plants. Under the influence of post-diagenetic oxidation in mineralized wood samples distribution of *n*-alkanes is changing. Diaster-13(17)-enes with 28 and 29 carbon atoms in molecule are present in the wood samples, while those with 29 atoms strongly prevail.

Keywords: fossil wood, Middle Jurassic, clays, organic matter, biomarkers, GC-MS

Reports according to the presence of fossil wood from the Middle Jurassic clays are known for many years (e.g., Jachowicz & Dybova-Jachowicz 1994, see Jachowicz & Dybova-Jachowicz 2003), but only recently this problem was studied in detail using paleobatanical and geochemical methods.

The mineralogical characteristics of fossil wood from Częstochowa region was made by Gut – Kałamaga (2000) in her master's thesis and it has demonstrated its phosphatic – carbonate mineralization.

Taxonomic research which was made by Philippe et al., (2006) has allowed to identify four genus of wood from that region, i.e.: Agathoxylon, Protopodocarpoxylon, Protaxodioxylon, Xenoxylon and a species *Xenoxylon phyllocladoides* GOTHAN.

Marynowski et al., (2007a, b; 2008) have carried out molecular research of wood from genus Protopodocarpoxylon ECKHOLD and *Xenoxylon phyllocladoides* GOTHAN, as well as taxonomic non-defined fragments of wood, showing the presence of such biomolecules as ferruginol, sugiol and other organic compounds including derivatives of the identified biomolecules. What also interesting, above mentioned compounds are the oldest described biomolecules detected in a sedimentary organic matter (OM) so far.

The aim of this work was to carry out a comprehensive bulk geochemical characteristic on numerous wood samples which has allowed to compare the contents of the organic fractions and their general molecular composition.

The study of content of total organic carbon (TOC) and total sulphur (TS) was made in order to demonstrate the differences between TOC values in respect of the degree of mineralization and the oxidation level. The thermal maturity of fossil wood was defined based on distribution of organic compounds.

MATERIALS AND METHODS

MATERIALS

Samples of Middle Jurassic wood occur in clays from the Polish Jura area. Wood samples come from active clay pits which are located in Czestochowa region: Gnaszyn Wienerberger, Sowa, Anna and near to that city between Grodzisko and Wreczyca Wielka (Grodzisko).



Middle Jurassic wood from Czestochowa region is characterized by diverse state of preservation and the type and degree of mineralization (Gut – Kałamaga 2000).

Due to the degree of mineralization wood can be divided into:

- not-mineralized wood fragments (xylite type) and gelified wood (jet type),
- wood fragments partly mineralized,
- > wood fragments completely mineralized.

Due to the degree of oxidation the following wood types can be distinguished:

- not-oxidized wood fragments,
- partly oxidized wood fragments,
- completely oxidized wood fragments.

METHODS

Instrumental geochemical analysis

тос

The total organic carbon (TOC) and total sulfur (TS) contentwere determined using Eltra Elemental Analyser model CS530.

Extraction and separation

Samples were Soxhlet-extracted with dichloromethane (DCM) for 48 h in cellulose thimbles. Extracts were further separated using silica gel TLC plates (Merck, 20 x 20 x 0.25 cm). Prior to separation, these plates were activated at 120°C for 1 h. Plates were then loaded with the DCM and developed with *n*-hexane. Aliphatic, aromatic and polar fractions were eluted and separated with DCM. The aliphatic fractions of sample were further analysed in details by gas chromatography-mass spectrometry (GC–MS).

GC-MS

The GC–MS analyses were performed with an Agilent 6890 Series Gas Chromatograph interfaced to an Agilent 5973 Network Mass Selective Detector and Agilent 7683 Series Injector (Agilent Technologies, Palo Alto, CA).

The detailed description of GC-MS method was made by Marynowski & Zatoń, (2010).

RESULTS

Organic matter fractionation

The test samples were separated into three organic fractions (aliphatic, aromatic, polar). Then, aliphatic fraction was analised using gas chromatography – mass spectrometry (GC-MS). Table 1 and triangular diagram (Fig.1) present the percentage distribution of each fraction content in tested wood samples.



Fig. 1. Triangular diagram showing percentage content of aliphatic (Al), aromatic (Ar) and polar fraction (Pol) in investigated wood samples.

Table 1. Percentage content of total organic carbon (TOC) and carbonates, mass and percentage content of extract, content of extract per gram of organic carbon (EOM) and percentage content of aliphatic (Al), aromatic (Ar) and polar fraction (Pol) in wood samples.

Sample	TOC	Carbonate	Extract	Extract % EOM				Fractions				
	[%]	[%]	mass [g]	extract	mg/gTOC	Al	Ar	Pol				
			Se	content		[%]	[%]	[%]				
			30	wa								
SOW	3.65	21.49	0.0306	0.0582	15.93	25	13	62				
JS1	1.65	28.08	0.0207	0.0726	15 (0)	17	22	(0)				
152	4.05	28.98	0.0397	0.0726	15.60	17	23	00				
SOW	5.68	41.34	0.0841	0.1430	25.17	1	8	91				
JS3								202/42				
SOW	5.87	9.00	0.0665	0.1120	19.08	2	5	93				
JS4 SOW	4.01	30.62	0 2328	0.4635	115.66	37	10	53				
JS6	1.01	50.02	0.2020	0.1055	110.00	21	10	55				
SOW	2.23	44.85	0.0226	0.0479	21.49	19	36	45				
JS7	2.47	20.72	0.0466	0.0001	26.52	0	~	0.5				
JS9	3.4/	38.72	0.0466	0.0921	26.55	9	0	85				
SOW	6.47	35.45	0.3149	0.6285	97.17	17	19	64				
J10												
SOW	3.95	41.68	0.1134	0.2243	56.76	3	9	88				
SOW	8.36	10.16	0.0623	0.1384	16.56	-	-	-				
J14												
			Gna	szyn								
GN JS2	7.49	12.05	0.0629	0.1324	17.68	13	16	71				
GN JS4	9.14	9.63	0.0677	0.1651	18.06	18	18	64				
GN JS5	68.5	-	0.3250	1.7857	26.07	14	22	64				
GN JS6	2.21	66.71	0.0410	0.0784	35.45	13	28	59				
GNU SR	2.01	85.74	0.0295	0.0587	29.26	17	18	65				
GNU SP	1.28	76.12	0.0123	0.0273	21.25	11	53	36				
GNU	1.06	57.38	0.0108	0.0238	22.44	50	14	36				
GN W20	41.60	0.37	0.0015	0.0033	0.08	12	35	53				
GN W20S	2.87	73.63	0.0205	0.0391	13.64	16	22	62				
11203			Grod	zisko								
GRO JS1	10.45	42.53	0.0948	0.1806	17.29	13	20	67				
GRO JS4	4.27	41.81	0.0107	0.0222	5.20	-	-	-				
001			Ar	ina								
ANNA J1	2.25	68.10	0.2328	0.4637	206.01	29	14	57				

Total organic carbon (TOC)

TOC values in the test samples are in wide interval from 1.06% to 68.50%. The largest amount of TOC have been found in non-mineralized or partly mineralized wood samples (GNW20, GNJS5; Fig.2).



Fig. 2. A comparison of percentage content of the TOC for Middle Jurassic clays (from Marynowski *et al.* 2007a) and fossil wood tested in this work, indicating higher values for notoxidized wood samples.

These two samples representing only the outer, carbonized part of wood (GNW20) or are rare, small pieces of non-mineralized wood with a poorly preserved structure (GNJS5). For comparison, sample GNW20S with TOC = 2.87% is the internal, mineralized part of the GNW20 wood fragment (Tab.1).

Carbonate content

Carbonate content is in the wide range of almost 1% CaCO₃ to 85% CaCO₃. What interesting, these samples which are poor in carbonates are in some cases characterized by low amounts of TOC. It would indicate the presence of other than the carbonate mineral phases from which most likely are phosphates and sulfides (pyrite, marcasite and sphalerite)



Fig. 3. Percentage content of carbonates in choosen non-oxidized wood samples.

Total sulphur (TS)

For some part of samples the content of total sulphur was determined (Fig.4). Obtained percentage values of total sulphur are very varied and do not show convergence with any other data (e.g, TOC, contents of carbonates). It suggests other than organic origin of sulphur in analysed fossil wood, what raises a presumption that its main sources are sulphides, especially pyrite and marcasite.



Fig. 4. A comparison of percentage content of total sulphur from selected samples of Middle Jurassic wood, analyzed using ELTRA CS 530.

n-Alkanes, isoprenoids

N-alkanes are hydrocarbons which characterized by a linear arrangement of carbon atoms in the structure. The most often, structure of *n*-alkanes is dependent on source of organic matter. Generally, short-chain *n*-alkanes (from C_{12} to C_{23}) are characteristic for algae's lipid fraction, some bacteria, as well as cyanobacteria. Moreover, these alkanes are characteristic components of trees trunks. In case of long-chain *n*-alkanes (from C_{25} to C_{33}) they are the derivatives of fatty acids present in waxes of higher plants. Therefore, the preponderance of long-chain alkanes over short-chain is usually characteristic for epicontinental sediments, rocks of terrestrial origin, as well as deltaic (Peters & Moldowan 1993), that means for facies rich in organic matter comes from higher plants.

There was elaborated row of indicators based on distributions of n- alkanes and isoprenoids. It has allowed to define organic matter parameters. Two parameters, present in sediments: pristane (Pr) and phytane (Ph) were used for this work. Isoprenoids are hydrocarbon compounds from polymerized isoprene units. Pristane and phytane are the branched acyclic hydrocarbons which have 19 and 20 carbon atoms in molecule (Peters & Moldowan 1993).

Unlike the Middle Jurassic clay samples, where long-chain and short-chain *n*-alkanes occur in similar concentrations (Marynowski *et al.* 2007a), in wood samples always outweigh short-chain *n*-alkanes present in the range from 15 to 23 carbon atoms in molecule (Table 2, Table 3). In general, the dominant *n*-alkanes are C_{17} , C_{18}

or C_{19} (Table 2, Fig 5), although in some samples there was an advantage of C_{23} *n*-alkane (Table 3).

 Table 2. Summary of indicators from *n*-alkanes and isoprenoids group.

(25- 31) C17 C18 (nC27+nC28+nC29) SOW 2.20 1.49 0.70 0.53 0.61 5.05 JS1	Sample	CPI	CPI	Pr/Ph	Pr/n-	Ph/n-	(nC17+nC18+nC19)/
Sow 2.20 1.49 0.70 0.53 0.61 5.05 JS1 SOW 1.70 1.42 0.60 0.65 0.62 3.71 SOW 1.70 1.42 0.60 0.65 0.62 3.71 SOW 1.76 2.30 0.49 0.64 0.70 4.01			(25-		C17	C18	(<i>n</i> C27+nC28+nC29)
Sowa SOW 2.20 1.49 0.70 0.53 0.61 5.05 JS1 SOW 1.70 1.42 0.60 0.65 0.62 3.71 SOW 1.70 1.42 0.60 0.64 0.70 4.01 SOW 1.76 2.30 0.49 0.64 0.70 4.01			31)				
SOW 2.20 1.49 0.70 0.53 0.61 5.05 JS1 SOW 1.70 1.42 0.60 0.65 0.62 3.71 JS2 .					Sowa		
JS1 SOW 1.70 1.42 0.60 0.65 0.62 3.71 JS2 SOW 1.76 2.30 0.49 0.64 0.70 4.01 JS3	SOW	2.20	1.49	0.70	0.53	0.61	5.05
SOW 1.70 1.42 0.60 0.65 0.62 3.71 JS2 SOW 1.76 2.30 0.49 0.64 0.70 4.01 JS3 SOW 1.29 1.57 0.60 0.78 0.69 3.09 JS4 SOW 2.28 2.48 0.79 0.82 0.88 19.19 JS6 SOW 1.81 1.55 0.79 0.85 0.81 10.93	JS1						
JS2 SOW 1.76 2.30 0.49 0.64 0.70 4.01 JS3	SOW	1.70	1.42	0.60	0.65	0.62	3.71
SOW 1.76 2.30 0.49 0.64 0.70 4.01 JS3 SOW 1.29 1.57 0.60 0.78 0.69 3.09 JS4 SOW 2.28 2.48 0.79 0.82 0.88 19.19 JS6 SOW 1.81 1.55 0.79 0.85 0.81 10.93	JS2						
JS3 SOW 1.29 1.57 0.60 0.78 0.69 3.09 JS4 SOW 2.28 2.48 0.79 0.82 0.88 19.19 JS6 SOW 1.81 1.55 0.79 0.85 0.81 10.93	sow	1.76	2.30	0.49	0.64	0.70	4.01
SOW 1.29 1.57 0.60 0.78 0.69 3.09 JS4 SOW 2.28 2.48 0.79 0.82 0.88 19.19 JS6 SOW 1.81 1.55 0.79 0.85 0.81 10.93	JS3	1.20	1.67	0.00	0.70	0.00	2.00
SOW 2.28 2.48 0.79 0.82 0.88 19.19 JS6 SOW 1.81 1.55 0.79 0.85 0.81 10.93	SOW ISA	1.29	1.57	0.60	0.78	0.69	3.09
JS6 SOW 1.81 1.55 0.79 0.85 0.81 10.93	SOW	2.28	2 48	0.79	0.82	0.88	10 10
SOW 1.81 1.55 0.79 0.85 0.81 10.93	156	2.20	2.40	0.19	0.02	0.00	17.17
	SOW	1.81	1.55	0.79	0.85	0.81	10.93
JS7	JS7						
SOW 1.66 1.76 0.51 0.73 0.79 7.82	sow	1.66	1.76	0.51	0.73	0.79	7.82
JS9	JS9						
SOW 1.51 1.84 0.47 0.67 0.68 5.17	sow	1.51	1.84	0.47	0.67	0.68	5.17
JS10	JS10						
SOW 1.29 1.09 0.55 0.69 0.79 1.63	SOW	1.29	1.09	0.55	0.69	0.79	1.63
JS13	JS13						
Gnaszyn				C	snaszyn		
GN JS2 3.73 1.89 0.89 0.92 0.71 3.63	GN JS2	3.73	1.89	0.89	0.92	0.71	3.63
GN JS4 1.42 1.23 0.98 1.17 0.78 2.98	GN JS4	1.42	1.23	0.98	1.17	0.78	2.98
GN JS5 0.98 0.98 0.37 1.55 1.39 1.85	GN JS5	0.98	0.98	0.37	1.55	1.39	1.85
GN JS6 2.30 1.41 0.48 0.99 0.83 2.57	GN JS6	2.30	1.41	0.48	0.99	0.83	2.57
GNU SR 1.58 1.38 0.70 0.75 0.71 3.85	GNU SR	1.58	1.38	0.70	0.75	0.71	3.85
GNU UT 1.78 1.97 0.62 0.94 1.08 52.58	GNU UT	1.78	1.97	0.62	0.94	1.08	52.58
GN W20 2.47 - 0.69 0.72 0.90 -	GN W20	2.47	-	0.69	0.72	0.90	-
GN 1.94 1.99 0.73 0.67 0.84 50.88	GN	1.94	1.99	0.73	0.67	0.84	50.88
W20S	W20S						
Grodzisko				G	rodzisko		
GRO 1.53 1.18 0.89 0.86 0.70 2.51	GRO	1.53	1.18	0.89	0.86	0.70	2.51
JS1	JS1						
Anna					Anna		
ANNA 1.87 1.36 0.94 0.92 0.88 9.81	ANNA	1.87	1.36	0.94	0.92	0.88	9.81
J1	J1						

Explanations to table 2:

Pr/nC₁₇ –pristane to *n*-heptadekane ratio (Peters & Moldowan 1993) Ph/nC₁₈ – phytan to *n*-oktadekane ratio (Peters & Moldowan 1993) CPI – Carbon Preference Index

$$CPI_{(25-31)} = \frac{(C_{25} + C_{27} + C_{29}) + (C_{27} + C_{29} + C_{31})}{[2(C_{26} + C_{28} + C_{30})]}$$

(Kotarba *et.al.* 1994)

$$CPI_{(Total)} = \frac{\left(C_{17} + C_{19} \dots + C_{27} + C_{29}\right) + \left(C_{19} + C_{21} \dots + C_{29} + C_{31}\right)}{\left[2\left(C_{18} + C_{20} \dots + C_{28} + C_{30}\right)\right]}$$

(Kotarba et.al. 1994)

Pr/Ph – pristane to phytane ratio (Peters & Moldowan 1993)

 $(nC_{17}+nC_{18}+nC_{19})/(nC_{27}+nC_{28}+nC_{29})$ – ratio of short-chain *n*-alkanes with 17, 18 and 19 carbon atoms in the molecule to long-chain *n*-alkanes with 28 and 29 carbon atoms in the molecule.

However, the values of the CPI (Table 2) are generally higher than 1 (an except for the sample GN JS5), which indicates the contribution of organic matter derived from higher plant waxes, which are characteristic of needles from gymnosperm plants, as well as leaves and epithelia etc. angiosperms (Peters *et al.*, 2005).

In case of analyzed samples the prevalence of longchain odd-over-even n-alkanes (Fig. 5) may resulting either from partial contamination of wood surrounding by clays, or the presence of a small amount of waxes in the tissues of wood, or the presence of both factors described above.

Diasterenes

Diaster-13(17)-enes which have 28 and 29 carbon atoms in molecule are present in the investigated samples, with preponderance of that with 29 atoms in the molecule (Tab. 4, Fig. 6; Fig. 7). Only in few wood samples the occurrence of diaster-13(17)-enes with 27 carbon atoms in the molecule have been detected, always in low concentration.

In distinction from wood samples, clays are characterized by presence of diaster-13(17)-enes which have 27 carbon atoms (Fig. 7). Moreover, diaster-13(17)-enes which have 30 carbon atomes are present in clays but its concentrations is quite low (Marynowski *et al.* 2007a). The presence of diaster-13(17)-enes with 27 carbon atoms in some wood samples suggests their contamination by clays.

DISCUSSION

Organic matter fractionation

The high concentrations of polar fraction is characteristic for organic matter of low thermal maturity (Tissot &Welte 1984). On the other hand, high content of aliphatic fraction may suggest the contribution of organic material typical of higher plants (cuticle waxes, etc.). Low content of aromatic fractions may be related to the oxidation of organic matter, resulting in formation of polar compounds like quinones or ketones (Wilkes *et al.* 1998).

The results of fractionation show a clear advantage of polar fraction in the tested extracts. Similarly, the high content of polar fraction was obtained in the work Marynowski *et al.* (2007a). Molecular investigation and vitrinite reflectance measurements have showed that the organic matter present in investigated fossil wood is at a very low stage of thermal transformation as indicated by: the presence of biomolecules (Marynowski *et al.* 2007 a and b) and the vitrinite reflectance of about 0.3% (Marynowski *et al.* 2007b). Samples which were an exception and containing a relatively high amounts of aromatic or aliphatic fraction are oxidized wood fragments from Gnaszyn. It follows that the oxidation processes of organic matter of terrestrial origin cause the percentage changes of organic fraction consisting of the increase in content of aromatic and aliphatic fractions for the polar fraction.

Total organic carbon (TOC)

Total organic carbon (TOC) is normally used to determine the organic carbon content in the samples of sedimentary rocks (Tissot & Welte 1984).

Previous studies indicate a general upward trend of total organic carbon content, extending with an increase in anoxic conditions in the water column during sedimentation of organic matter (eg, Demaison & Moore 1980).

However, Calvert *et al.* (1996) are proving that high content of organic carbon may also be formed under anoxic conditions of sedimentation, and the main factor determining its amount in sediment is called organic productivity (the high activity of microorganisms – especially the algae).

Most of the samples demonstrate TOC values in the range of 2% - 10% (Table 1, Fig 2). These values are consistent with the values obtained for seven wood samples from the location Kawodrza, Blanowice, Żarki, Anna, Gnaszyn (Marynowski *et al.* 2007a). These authors have obtained TOC values in the range of 7% - 14% for mineralized and partly mineralized wood and 30% - 40% for poorly mineralized wood.

n-Alkanes, isoprenoids

Pristane to phytane ratio (Pr / Ph) is mainly used to reconstruct the sedimentary conditions of organic matter (Didyk *et al.* 1978). The wood values of this ratio is less than 1 (Table 2), what may suggest rather reductive conditions prevailing in the basin.

However, due to the specific nature of the samples, which are pieces of wood representing a fairly uniform type of organic matter, the above described ratio does not reflect the conditions during sedimentation of the investigated samples.

The indicators Ph/n-C18 and Pr/n-C17 are usually less than 1, indicating a higher content of *n*-alkanes relative to isoprenoids in wood samples. In two cases (GNJS4, GNJS5) slightly higher values may be related to its partial biodegradation. During the biodegradation process short-chain *n*-alkanes are removed at the first stage (Connan 1984).

An interesting aspect is to compare the distribution of *n*-alkanes in one sample of wood, with oxidized zone, however it is difficult to determine whether it is diagenetic oxidation, or - more likely - post-diagenetic oxidation resulting from weathering. Oxidized zone (GNUSR) is characterized by the dominant of *n*-alkanes C_{18} and C_{19} , as well as the occurrence of relatively high concentrations of *n*-alkanes C_{22}

and C₂₃ and also the relatively intensive *n*-alkanes rang-

ing from C_{26} to C_{33} . Besides, in the sample GNUSR the predominance of long-chain odd-over-even *n*-alkanes (C₂₇, C₂₉, C₃₁, C₃₃) over even *n*-alkanes (C₂₆, C₂₈, C₃₀, C₃₂) was observed. The transition zone and the oxidized zone in samples have a very similar distribution of *n*-alkanes, where dominated short-chain n-alkanes and long-chain n-alkanes are present in very small quantities. This is particularly visible in the oxidized sample. The oxidation of organic matter, during which a reduction of long-chain *n*-alkanes occur is known from literature data with respect to weathering (oxidation) of clays (Elie et al. 2000), as well as the lower- and upper Triassic clays from southern Poland (Marynowski & Wyszomirski 2008). However oxidation processes could be strongly connected with organic matter biodegradation.

Diasterenes

Steroids occurring in the sedimentary organic matter have formed from sterols and stanols, biosynthesized by living organisms through the transformation taking place at an early stage of organic matter burial.

Steranes and diasteranes - the final products of diagenetic transformations of sterols and stanols - are common compounds, described in the coals, crude oils and other sedimentary rocks. Steranes usually contain from 27 to 29 carbon atoms in the structure, although in specific sedimentary environments also steranes with 21, 22, 26 and 30 carbon atoms in the molecule can be present.

The formation of diaster-13(17)-enes is possible only when the organic matter "neighbors" with clay minerals during diagenesis. It causes the rearrangement of carbon skeleton and the formation of diaster-13(17)-enes and in much lower concentrations spirosterenes.

During the subsequent diagenetic transformation processes (mainly by affecting the temperature ingeological time), from diasterenes arise diasteranes, commonly found in sedimentary organic matter (Peakman *et al.* 1987). Diasterenes are precursors of common in sedimentary organic matter, diasteranes and are found in thermally immature samples (Rubinstein *et al.* 1975).

According to literature data (Brassell *et al.* 1985) C_{29} -diaster-13(17)-enes come from higher plants, while C_{27} -diaster-13(17)-enes are from marine organic matter, mainly algae. In Middle Jurassic clays both marine and terrestrial diasterenes occur and these second are dominated (Marynowski *et al.* 2007a). However in fossil wood almost only C29- diaster-13(17)-enes are present, while other diasterenes are contamination of wood samples by clays (Figs. 5 & 6, Table 4).

Sample	[%]	[%]	[%]	[%]	[%]	[%]	[%]	[%]	[%]	[%]	[%]	[%]	[%]	C28 [%]	[%]	[%]	[%]	[%]	[%]
1000000000000	1983.0	02020	7559451	10000	12122	1200	0002	1000	Sowa	2220	10000	32322	282.2	10020	6353	1000	22025	8128	1240-255
SOWJS1	0.5	2.7	8.8	10.8	9.5	7.1	14.7	8.0	17.1	5.3	5.4	2.7	2.6	1.5	1.6	0.7	0.9	0.0	0.0
SOWJS2	0.0	1.2	5.6	9.9	9.3	7.6	13.3	9.8	14.7	7.3	7.4	4.0	2.9	1.7	2.1	1.2	2.2	0.0	0.0
SOWJS3	0.0	1.3	4.3	8.0	10.0	10.2	11.5	12.9	20.7	4.8	5.6	2.0	1.9	1.1	2.6	0.7	2.5	0.0	0.0
SOWJS4	0.0	1.3	6.6	12.3	12.8	9.5	8.9	12.9	6.2	4.6	3.6	3.2	3.4	2.5	4.3	1.7	4.0	1.0	1.2
SOWJS6	0.0	9.3	13.3	15.6	13.5	9.9	10.5	б.б	11.3	2.7	3.0	0.8	0.9	0.4	1.0	0.4	1.0	0.0	0.0
SOWJS7	4.7	7.5	13.0	17.2	12.7	8.5	9.5	7.1	6.1	3.7	2.5	1.8	1.7	0.8	1.4	0.5	1.1	0.0	0.0
SOWJS9	0.0	2.0	12.0	22.0	15.0	10.6	7.7	7.2	5.2	3.6	2.2	1.7	2.4	1.3	2.5	1.3	3.1	0.0	0.0
SOWJS10	0.0	1.1	9.2	19.7	15.8	11.5	8.8	9.3	5.6	4.4	3.0	2.4	1.7	3.2	1.1	3.3	0.0	0.0	0.0
SOWJS13	0.0	1.5	7.8	12.4	9.0	6.1	5.1	5.9	6.0	6.9	7.2	7.9	7.2	5.6	5.1	2.7	3.5	0.0	0.0
								C	Jnaszy	m									
GN JS2	0.0	1.0	3.7	5.4	6.0	4.3	14.7	5.0	31.0	4.1	14.4	1.5	2.0	1.0	1.2	3.3	1.5	0.0	0.0
GN JS4	0.0	2.3	6.9	10.6	12.1	9.0	8.7	7.7	5.8	4.7	4.6	3.7	4.2	3.0	4.9	8.3	3.5	0.0	0.0
GN JS5	0.0	0.7	2.6	7.9	6.8	7.4	8.7	14.0	10.3	11.1	9.2	7.7	4.2	2.9	2.3	2.0	2.4	0.0	0.0
GN JS6	0.0	0.6	3.6	8.8	9.1	7.3	11.3	8.6	26.9	3.7	4.6	2.1	2.7	1.8	3.8	3.1	1.8	0.0	0.0
GNU SR	0.0	3.0	8.6	12.9	10.6	7.9	7.9	8.9	11.8	5.9	5.2	3.2	3.2	2.3	2.9	1.6	2.2	1.0	0.9
GNU UT	0.0	5.8	12.7	18.0	17.4	13.8	11.0	7.6	7.4	2.7	1.7	0.6	0.4	0.2	0.3	0.1	0.3	0.0	0.0
GN W20	0.0	9.3	18.0	21.3	14.0	10.1	13.3	5.4	5.1	1.9	1.6	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0
GN W20S	0.0	9.9	16.6	18.2	14.2	10.3	10.3	7.8	6.5	2.6	1.7	0.9	1.0	0.0	0.0	0.0	0.0	0.0	0.0
								G	rodzis	ko									
GRO JS1	0.0	2.0	10.3	14.1	10.4	7.0	6.1	6.2	6.8	5.2	8.4	3.7	8.8	3.0	2.0	6.0	0.0	0.0	0.0
									Anna										
ANNA JI	0.0	11.6	15.9	17.6	11.3	8.2	8.8	5.7	4.6	3.1	2.8	2.0	1.8	1.3	1.5	0.9	2.0	0.9	0.0

Table 3. Proportional content of *n*-alkanes counted based on surface area of the peaks.

Table 4. Proportional participation of diaster-13(17)-enes from 27 to 29 carbon atoms in molecules.

Sample	27(\$+R)	28(S+R)	29(S+R)		
	[%]	[%]	[%]		
	Som	a			
SOW JS1	0.0	8.1	91.9		
SOW JS2	0.0	4.9	95.1		
SOW JS3	2.8	5.5	91.7		
SOW JS4	0.0	9.3	90.7		
SOW JS6	0.0	6.9	93.1		
	Gnas	zyn			
GN J S2	1.2	3.9	94.8		
GN JS4	2.4	11.1	86.5		
GN J S5	0.0	21.2	78.8		
GN J S6	0.8	3.4	95.8		
GNU SR	0.0	12.3	87.7		
GNU UT	0.0	4.7	95.3		
GN W20	0.0	5.8	94.2		
GN W20S	0.0	7.5	92.5		
	Ann	1 a			
ANNA J1	0.0	2.4	97.6		



Fig. 5. Comparision of proportional content of diasterenes in (a) clays, wood and carbonate concretions from Marynowski *et al.* (2007a) and (b) in fossil wood analysis from this work.



Fig. 6. Examples of diaster-13(17)-enes distribution characteristic for the Middle Jurassic deposits including clay sample (from Marynowski et. al., 2007a) as well as wood samples from clay pit Sowa, Anna and Gnaszyn. Shortcuts explanations in Table 1.

CONCLUSIONS

- Macroscopic observations, microscopic studies and literature data revealed that Middle Jurassic wood from Częstochowa area has a different state of preservation, and various types and degrees of mineralization and oxidation.
- Obtained results of organic matter fractionation illustrate a clear domination of polar fraction in the investigated extracts. Dominant content of polar fraction are characteristic of the samples in the low thermal maturity stage. Oxidized wood fragments from Gnaszyn GNUSP and GNUUT are an exception and contain a relatively high amounts of aromatic and aliphatic fraction. Oxidation (weathering) of organic matter of terrestrial origin changing the percentage consisting of organic fraction, it means growing content of aromatic and aliphatic fraction.
- TOC values in tested samples are in a wide range from 1.06% to 68.50%. The greatest amount of total organic carbon were in the non-mineralized or poorly mineralized wood. Most of the samples are mineralized wood fragments, showing the TOC values in the range of 2% - 10%.
- Percentage content of carbonate in experienced wood samples constitute in a wide range from less than 1% CaCO₃, to above 85% CaCO₃. An exception are poorly mineralized samples where carbonate content ranges from 50% - 90%.
- The percentage of the total sulfur content is very varied for wood samples and do not show convergence with other data such as TOC, carbonate content, etc.
- Unlike the Middle Jurassic clay samples, where long-chain and short-chain *n*-alkanes occur in similar concentrations in wood samples always shortchain *n*-alkanes from 15 to 23 carbon atoms in molecule dominating.
- The values of the CPI are generally higher than 1 which indicates the contribution of organic matter derived from higher plant waxes, which are characteristic of needles from gymnosperm plants. In case of analyzed samples the prevalence of long-chain *n*-alkanes odd-over-even may result either from partial contamination of wood surrounding by clays, or the presence of a small amount of waxes in the tissues of wood, possibly is also the presence of both factors described above.
- Under the influence of post-diagenetic oxidation in mineralized wood samples distribution of *n*-alkanes is changing.
- In tested wood samples are diaster-13(17)-enes with 28 and 29 carbon atoms in molecule, while those with 29 atoms strongly prevail. Diaster-13(17)-enes are present only in part of the wood samples, what may result from the impact of diagenetic processes on the presence or absence of these compounds in wood samples.

REFERENCES

- Brassell S.C., Eglinton G., Mo, F. J. (1985) Biological marker compounds as indicators of the depositional history of the Maoming oil shale. *Organic Geochemistry* 10, 927-941
- Calvert S. E., Bustin R. M. (1996) Ingall Influence of water column anoxia and sediment supply on the burial and preservation of organic carbon in marine shales. *Geochimica et Cosmochimica Acta* 60, 1577-1593.
- Connan J. (1984) Biodegradation of crude oils in reservoirs. In: Brooks J., Welte D. (Eds.). Advances in Petroleum Geochemistry, vol. 1. Academic Press, London, pp. 299-335.
- Demaison, G.J. and Moore, G.T. (1980) Anoxic environments and oil source bed genesis. *Am. Assoc. Pet. Geol. Bull.* 64, 1179-1209.
- Didyk B.M., Simoneit B.R.T., Brassell S.C., Eglinton G. (1978) Organic geochemical indicators of palaeoenvironmental conditions of sedimentation. *Nature* 272, 216–222.
- Elie, M., Faure, P., Michels, R., Landais, P., Griffault, L. (2000) Natural and laboratory oxidation of low-organic-carboncontent sediments: comparison of chemical changes in hydrocarbons. *Energy & Fuels* 14, 854-861.
- Gut–Kałamaga M. (2000) Mineralizacja drewna Jury środkowej z rejonu Częstochowy. Praca magisterska. Archiwum Uniwersytetu Śląskiego, Sosnowiec (in Polish).
- Jachowicz J., Dybova Jachowicz S. (2003) Paleobotanika. Wydawnictwo Uniwersytetu Śląskiego, Katowice, 1–210 (in Polish).
- Kotarba M., Kowalski A., Więcław D. (1994) Nowa metoda obliczeń wskaźnika CPI i wykorzystanie dystrybucji nalkanów i izoprenoidów prospekcji naftowej – instrukcja metodyczna analiz z zakresu geochemii organicznej dla potrzeb poszukiwań naftowych. Materiały wewnętrzne AGH: 82-91 (in Polish).
- Marynowski L., Zatoń M., Simoneit B.R.T., Otto A., Jędrysek M.O., Grelowski C., Kurkiewicz S. (2007a) Compositions, sources and depositional environments of organic matter from the Middle Jurassic clays of Poland. *Applied Geochemistry* 22, 2456-2485.
- Marynowski L., Otto A., Zatoń M., Philippe M., Simoneit B.R.T. (2007b) Biomolecules preserved in 168 million year old fossil conifer wood. *Naturwissenschaften* 94, 228-236.
- Marynowski L., Wyszomirski P. (2008) Organic geochemical evidences of early diagenetic oxidation of the terrestrial organic matter during the Triassic arid and semi arid climatic conditions. *Applied Geochemistry* 23, 2612-2618
- Marynowski L., Philippe M., Zatoń M., Hautevelle Y. (2008) Systematic relationships of the Mesozoic wood genus Xenoxylon: an integrative biomolecular and palaeobotanical approach. *N. Jb. Geol. Paläont. Abd.* 247, 177-189.
- Marynowski L., Zatoń M. (2010) Organic matter from the Callovian (Middle Jurassic) deposits of Lithuania: Compositions, sources and depositional environments. *Applied Geochemistry* 25, 933–946.
- Peakman J., Torren M., Maxwell J. R. (1987) Early diagenetic pathways of steroids alkanes. *Adwances in Organic Geochemistry* 13, 583-592.
- Peters K.E., Moldowan J.M. (1993) The biomarkers guide. Interpreting molecular Fossils in Petroleum and Ancient Sediments: 363p. Prentice Hall, Englewood Cliffs, New Jersey.
- Philippe M., Barbacka M., Gradinaru E., Iamendei E., Iamandei S., Kázmér M., Popa M., Szakmány G., Tchoumatchenco P., Zatoń M. (2006) Fossil wood and Mid-Eastern Europe

terrestrial palaeobiogeography during the Jurassic-Early Cretaceous interval. *Review of Palaeobotany and Paly-nology* 142, 15-32.

- Rubinstein I., Sieskind O., Albrecht P. (1975) Rearranged sterenes in a shale: Occurrence and simulated formation. *Journal of the Chemical Society*; Perkin Transactions I, 1833-1836.
- Tissot B., Welte D. H. (1984) Petroleum Formation and Occurence-Springer-Verlag, New York, s.699.
- Wilkes H., Disco U., Horsfield B. (1998) Aromatic aldehydes and ketones in the Posidonia shale, Hils syncline, Germany. Organic Geochemistry 29, 107-117.

Abstrakt

Makroskopowe obserwacje, mikroskopowe badania i dane literaturowe wykazały, że środkowojurajskie drewno z okolic Częstochowy ma różny stan zachowania oraz różne rodzaje i stopień mineralizacji oraz utlenienia. Uzyskane wyniki rozdziału frakcyjnego pokazują wyraźną przewagę frakcji polarnej w badanych ekstraktach. Wartości TOC dla badanych próbek wahają się w szerokim zakresie od 1,06% do 68,50 %. Największa ilość TOC występuje w nie zmineralizowanych lub słabo zmineralizowanych próbkach drewna. Większość próbek to zmineralizowane fragmenty drewna, pokazując wartości TOC w zakresie od 2% - 10%. Zawartość procentowa węglanów w badanych próbkach znajduje się w szerokim zakresie od mniej niż 1% CaCO₃ do ponad 85% CaCO₃. Procentowa zawartość siarki całkowitej jest bardzo zróżnicowana i nie wykazuje zbieżności z innymi danymi, takimi jak TOC czy zawartość węglanów, itp. W odróżnieniu od środkowojurajskich iłów, gdzie długołańcuchowe i krótkołańcuchowe n-alkany występują w podobnych stężeniach , w próbkach drewna zawsze przewyższają nalkanów krótkołańcuchowe, występujące w zakresie od 15 do 23 atomów węgla w cząsteczce. Wartości wskaźnika CPI są generalnie wyższe niż 1, co wskazuje na udział materii organicznej pochodzącej z wosków roślin wyższych. Pod wpływem post-diagenetycznego utleniania zmineralizowanych próbek drewna zmienia się dystrybucja n-alkanów. W badanych próbkach obecne są diaster-13 (17)-enes o 28 i 29 atomów wegla w cząsteczce, natomiast te o 29 atomów znacznie przeważają.

Słowa kluczowe: skamieniałe drewno, jura środkowa, iły, materia organiczna, biomarkery, GC-MS