An overview of lithotype associations of Miocene lignite seams exploited in Poland

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Abstract

Currently, three stratigraphically distinct lignite seams of Early to Middle Miocene age are exploited in Poland, namely the third Ścinawa lignite seam (ŚLS-3), the second Lusatian lignite seam (LLS-2) and the first Mid-Polish lignite seam (MPLS-1). All of these are composed of numerous macroscopically distinguishable layers defined as lignite lithotypes. In the present paper, the lithotypes of Polish lignites are grouped into seven major lithotype associations that originated in various types of mire. Therefore, an approximate reconstruction of mire type can be based on lignite lithotypes. Within the Polish lignite seams examined, the commonest in order of importance are: xylodetritic (XDL), detroxylitic (DXL), detritic (DL) and xylitic (XL) lithotype associations, mostly with a massive (m) or horizontal (h) structure. They are particularly dominant in lignite opencasts belonging to the Konin and Adamów mines. However, in the lowermost seams at the Turów and Bełchatów mines, a substantial part of the seams comprises the bitumen-rich (BL) lithotype association. These seams also lignite lithotypes that in large quantities have a gelified (g) and/or nodular (n) structure. In contrast, lignites from the Sieniawa mine are characterised by an admixture of the best-developed lithotype associations of both fusitic (FL) and weathered (WL) lignites. Moreover, the vast majority of these lignites have a folded (fo) and/or faulted (fa) structure, because they were completely deformed by glacitectonics.

Keywords: brown coal, Cenozoic, northern Europe, lithotypes

1. Introduction

Poland is one of the most significant lignite producers, ranking second in the European Union and fourth worldwide, with an annual extraction of about 63–66 Mt. Therefore, lignite plays a key role in electricity production, supplying 30–33% of Poland’s total electricity, that is, about 53–55 TWh (Pietraszewski, 2015).

Currently, Polish lignite production is conducted at four large mines and a single small one, in nine opencasts. The largest are the Adamów Lignite Mine with the Adamów and Koźmin N opencasts, the Bełchatów Lignite Mine with the Bełchatów and Szczerców opencasts, the Konin Lignite Mine with the Jóźwin IIB, Drzewce and Tomisławice opencasts and the Turów Lignite Mine with the Turów opencast. The smallest is the Sieniawa Lignite Mine with the Sieniawa opencast (Fig. 1).

Taking into account the size of lignite production and number of opencasts, it can be stated that the Polish lignite mines provide an excellent opportunity to examine lignite seams in detail. The seams exploited vary in age and, in some cases, belong to the thickest in Europe and around the world. These lignite seams also differ petrographically; that is, they consist of various lithotypes (e.g., Brzyski, 1981; Wagner, 1996; Kwiecienśka & Wagner, 1997; Majewski, 2002; Kasinski et al., 2010; Widera, 2007, 2012; Nowak, 2012; Bielowicz, 2013; Fabiańska & Kurkiewicz, 2013; Bielowicz & Kasinski, 2014).

The major aims of this paper are threefold: 1) to present a brief, lucid explanation of how to distinguish and describe lignite lithotypes within
a lignite seam in the field; 2) to provide examples of the most interesting lignite lithotypes from the exploited Polish lignite seams; and 3) to reconstruct the mire type based on the resultant lignite lithotypes. A secondary goal of this research is to discuss the relationship between lignite lithotypes and the stratigraphical position of the lignite seam as well as between lignite lithotypes and the location of the lignite-bearing area within Poland.

2. Geological setting

Lignite mining is concentrated in five areas of central and western Poland (Fig. 1). This is due to the fact that almost all lignite deposits of economic value are located in the same territories (e.g., Ciuk & Piwocki, 1990; Kasiński & Piwocki, 2002). Generally, the lignite-bearing areas cover approximately 70,000 km², that is, about one-third of the Poland’s total territory (Piwocki, 1992). On the other hand, the minable geological resources of 90 of the largest Polish lignite deposits were determined to be more than 23.5 Gt at the end of 2014 (Szuflicki et al., 2015). Most of Poland, inclusive of the lignite-bearing areas, belongs to the northwest European Paleogene–Neogene Basin (Vinken, 1988). Due to poverty of other palaeontological remains, lignite seams play a key stratigraphic role especially in the case of the Neogene deposits. This makes it possible to correlate the organic-rich deposits from western Poland with those from eastern Germany (compare Piwocki & Ziembińska-Tworzydło, 1997; Widera et al, 2008). Thus, the mined Polish seams, that is, the third Ścinawa lignite seam (ŚLS-3), the second Lusatian lignite seam (LLS-2), and the first Mid-Polish lignite seam (MPLS-1) match the German third, second, and first Lusatian lignite seams, respectively (compare Grimm et al., 2002; Widera, 2016a).

These lignite seams are of Early and Middle Miocene age (Fig. 2). They all cover areas that subsided tectonically at that time. In other words, the investigated lignite seams mostly fill grabens, which range in depth from a few dozen metres to more than 500 m (Kasiński, 2000; Wagner et al., 2000; Widera, 2007, 2016b; Kasiński et al., 2010; Widera & Haluszczyk, 2011). The oldest third Ścinawa lignite seam (ŚLS-3) is currently exploited in the Bełchatów, Szczebrzeszyn and Turów opencasts. The second Lusatian lignite seam (LLS-2) comprises the uppermost parts of the mined seams in the same opencasts, and additionally it is also exploited in the Sieniawa opencast mine. Conversely, the first Mid-Polish lignite seam (MPLS-1) is now excavated only in all of the opencasts belonging to the Adamów and Konin lignite mines (Figs. 1, 2; Widera, 2016a).

Polish lignite deposits are among the genetically most differentiated in the world (Ciuk, 1968;
Kasiński & Piwocki, 2002; Widera, 2016b). In some
cases, the lignite seams vary in thickness from a few
metres to more than 250 m (Piwocki, 1992; Widera,
2016b). All lignites exploited represent humic low
rank coal, i.e. ortholignite (Kwiecińska & Wagner,
1997, 2001). On the other hand, palynological data,
the relatively high mean ash yield (between 7.7 and
22.9%, calculated on a dry basis), and the lithotype
associations (described in detail in the present pa-
ter) provide evidence that the predominant part of
the Polish lignites formed in low-lying mires (e.g.,
Kasiński et al., 2010; Widera, 2012, 2016a; Mastej et
al., 2015).

3. Methods

Macroscopic observations and measurements of
lignite constituents have been made on two-dimen-
sional surfaces, that is, on lignite quarry surfaces.
It should be clearly stated that these actions must be
conducted only in the field on fresh lignite surfaces.
Thus, this method can be simply illustrated in three
steps (Fig. 3).

In the first step, on the basis of clearly visible
textural (position, size and number of xylites, dif-
fences in colour, mineral intercalations, etc.) and
structural features (massiveness, type of stratifica-
tion, deformations, etc.), boundaries of layers (L1,
L1, L3) are determined (Fig. 3A).

The second step is to measure the xylite length
within the lignite layer examined. In order to obtain
more representative data, it is proposed to make
the measurements, for example, along three parallel
lines (Fig. 3B). This is certainly the most important
step in determining the lithotype. Of course, the
difference between the total thickness of the layer an-
alysed and the total length of xylites gives the total
length of a detrital matrix.

In the third step, after adding the lignite struc-
ture, the lithotype may be appropriately named. In
this case, the so-called ‘10% rule’ is commonly used
by European lignite researchers. This was also rec-
ommended by the ICCP (International Committee
for Coal Petrology) in 1993 (compare Widera, 2012;
Flores, 2013). Finally, the lignite-lithotype codifica-
tion is suggested to simplify and shorten the full
description (Fig. 3C; Table 1).

The first version of lithotype codification, com-
prising both textural and structural features of lig-
nite, was proposed by the author (Widera, 2012). It
is corrected and supplemented in the present paper.
Thus, the texture is coded with capital letters and

<table>
<thead>
<tr>
<th>Lithotype texture</th>
<th>Code</th>
</tr>
</thead>
<tbody>
<tr>
<td>xylitic lignite</td>
<td>XL</td>
</tr>
<tr>
<td>detritic lignite</td>
<td>DL</td>
</tr>
<tr>
<td>detroxylitic lignite</td>
<td>DXL</td>
</tr>
<tr>
<td>xylodetritic lignite</td>
<td>XDL</td>
</tr>
<tr>
<td>bitumen-rich (bituminiferous)</td>
<td>BL</td>
</tr>
<tr>
<td>fusitic lignite</td>
<td>FL</td>
</tr>
<tr>
<td>weathered lignite</td>
<td>WL</td>
</tr>
</tbody>
</table>

Table 1. Codification of textural and structural features
of lignite lithotype associations (after Widera, 2012;
modified and supplemented).

<table>
<thead>
<tr>
<th>Lithotype structure</th>
<th>Code</th>
</tr>
</thead>
<tbody>
<tr>
<td>massive</td>
<td>m</td>
</tr>
<tr>
<td>horizontally stratified</td>
<td>h</td>
</tr>
<tr>
<td>deformed (fractured, folded, faulted)</td>
<td>d (fr, fo, fa)</td>
</tr>
<tr>
<td>gelified</td>
<td>g</td>
</tr>
<tr>
<td>nodular</td>
<td>n</td>
</tr>
</tbody>
</table>

the structure is coded with lower case letters (Table 1). For more information on the former Polish and international nomenclature and codification of lignite lithotypes reference is made, for example, to the following contributions: Suss & Sontag (1966), Sontag & Suss (1976), Brzyski (1981), Wolf (1988), Kwiecińska & Wagner (1997), Markič & Sachsenhofer (1997), Taylor et al. (1998), Kolcon & Sachsenhofer (1998, 1999), Ticleanu et al. (1999), Markič et al. (2001), Widera (2012) and Flores (2013). Detailed criteria for distinguishing lignite lithotypes and their best-developed examples among the currently exploited Polish lignite seams are discussed below.
3. Characteristics of lignite lithotype associations in Poland

3.1. Xylitic lignite association (XL)

3.1.1. Description
The xylitic lignite association is characterised by the presence of more than 90% ('10% rule') fossilised wood particles, that is, xylites. They must be macroscopically distinguishable woody fragments that are larger than 1 cm, such as stumps, trunks, branches, twigs and cones. They are usually light to dark brown in colour and their internal structure is easily visible (Figs. 3, 4). Obviously, the remainder of the lithotype forms the matrix, which consists of black, amorphous, fine-detrital organic matter and woody particles smaller than 1 cm in diameter (Kwiecińska & Wagner, 1997). The horizontal, gelified and deformed structures may characterise the xylitic lignites. However, the commonest and typical feature of this lithotype association is a massive structure (compare Figs. 3, 4).

3.1.2. Interpretation
Based on the pioneering research of Teichmüller (1958), when the lignite lithotype is known, an initial mire type can be roughly defined. Thus, the xylitic lignite association corresponds to a dry forest swamp as the original depositional environment (Teichmüller, 1958, 1989).

The presence of xylitic lignites within a seam, especially the so-called ‘stump horizons’, provides a convincing record of dry conditions in mire evolution when the ground water-table was relatively low (e.g. Ticleanu et al., 1999; Diessel et al., 2003, 2007). Such horizons with stump remains of *Taxodium*, *Juniperus* and *Glyptostrobus* are exceptionally well developed in the Turów opencast mine within ŚLS-3 and LLS-2 (Kasiński et al., 2010). In other opencasts, the xylites most often lie horizontally (trunks, branches, twigs), either in growth position (stumps, roots) or arranged chaotically (Figs. 3, 4). Xylitic lignites are occasionally found within Polish lignite seams, that is, forming up to 10% of seam thickness (Widera, 2014a, 2014b). This is due to the fact that the dry forest swamp is characterised by environmental conditions that are favourable for the decomposition of plant remains rather than their accumulation (Brzyski, 1981).

3.2. Detritic lignite association (DL)

3.2.1. Description
In contrast to the lithotype described above, the detritic lignite association contains at least 90% ('10% rule') plant detritus. It is predominantly of a dark grey to black colour (Figs. 3A, C; 4B). The remainder of this lignite association consists of other lithotypes. These are most often xylites, which are distributed irregularly within the detritic matrix. Therefore, a massive structure is the most characteristic of the detritic lignite association. Other structures, including a gelified one, are also very common in some cases (Figs. 5A, 5B). This specific structure also applies to the gelified xylites (Fig. 5C).

3.2.2. Interpretation
The detritic lignite association most often formed in mire subenvironments related to a fen, open water or a treeless reed marsh (Teichmüller, 1958, 1989). The parent organic material for the creation of this lithotype association could be aquatic plants, sedges and reed-like vegetation, respectively (Kolcon & Sachsenhofer, 1999; Ticleanu et al., 1999). However, woody vegetation cannot be excluded as the parent material for the detritic lignite, especially in the case of a high degree of plant decomposition.

Here, a gelified lignite should be briefly discussed, which is clearly visible in some parts of ŚLS-3 and LLS-2 at the Belchatów, Szczerców, and Turów opencast mines (Figs. 5A-C; Wagner et al., 1983; Kasiński, 2000; Kasiński et al., 2010). Effects of the gelification process rarely occur in other Polish lignite opencasts (Fig. 5D; Bechtel et al., 2007; Fabiańska & Kurkiewicz, 2013). The gelification can be produced by both geochemical and biochemical processes such as the decay of fossil wood. In the first case, the organo-mineral compounds are formed during the progressive doppleritisation of fossil wood (e.g., Brzyski, 1981; Wagner, 1982). In the latter case, the humic acids are transformed into purely organic compounds, that is humins, due to enhanced microbial activity (e.g., Wagner et al. 1983; Diessel, 1992; Bechtel et al., 2003, 2007). The result of these processes is an amorphous, black and intensively glassy substance; hence, it can be easily distinguished in the field (Fig. 5A-D).

3.3. Detroxylitic lignite (DXL) and xylodetritic lignite (XDL) associations

3.3.1. Description
Both associations consist of an admixture of xylites and fine-detrital matrix, amounting to a total content of more than 90% according to the ‘10% rule’ (Kwiecińska & Wagner, 1997; Kolcon & Sachsenhofer, 1998). In the case of detroxylitic lignite, xylites prevail over organic detritus. In contrast, xylodetritic lignite is in fact detritic lignite...
with a considerable amount of xylites (Brzyski, 1981; Widera, 2012). These lithotype associations may represent all the structures that are typical of the examined Polish lignites (Table 1).

### 3.3.2. Interpretation

These two lithotype associations represent transitional environmental conditions between dry forest swamp and open water. Thus, detroxylitic lignites are attributable to the wet forest swamp, whereas xylodetritic lignites are typical of the bush moor (e.g. Diessel, 1992; Kolcon & Sachsenhofer, 1999). According to the nomenclature proposed by Teichmüller (1958, 1989), they can also be called the *Taxodium–Nyssa* and *Myricaceae–Cyrillaceae* swamps, respectively. Moreover, the water table
had to change from a long-lasting and low level to a relatively high level with respect to the mire surface (e.g., Diessel et al., 2000).

Detroxylitic and xylodetritic lignites are the commonest lithotype associations among Polish lignites. In some cases they constitute more than 85% of the lignite’s seam thickness. This situation occurs in some sections of MPLS-1 in the opencasts belonging to the Adamów Lignite Mine (Nowak, 2012).

3.4. Bitumen-rich lignite association (BL)

3.4.1. Description
In the literature, the bitumen-rich lignite is also known as sapropelic coal, clayey coal or pyropisite (Ticleanu et al., 1999; Kasiński, 2000). It is also divided by Wagner (1996) into bituminiferous and semi-bituminiferous lignite. It has a characteristic colour that can vary from pale yellow to yellowish grey (Fig. 5A, 5B). Thus, it contrasts sharply with other lithotypes in vertical section. On the other hand, the bitumen-rich lignite layers extend laterally from a few dozen to a few hundred metres (Fig. 5E). Due to its origin, it is mostly characterised by a massive or horizontal structure. In the latter case, the lamination is often underlined by strongly gelified strings or lenses (Fig. 5B).

3.4.2. Interpretation
Petrographic features of the bitumen-rich (bituminiferous) lignite indicate that phytoplankton sedimentation took place in small and shallow lakes occurring on the mire surface (Wagner, 1996). However, some researchers believe that this lithotype association formed in the shallow, coastal zone of the vast lakes (e.g. Traverse, 1988; Kasiński et al., 2010). More organic matter could be deposited in shallower water (< 2 m), while sedimentation of mineral particles such as clays predominated in deeper water (> 2 m) (e.g. Smith et al., 1989; Diessel et al., 2000).

The bitumen-rich lignite is well developed within ŠLS-3 and LLS-2 in the Belchatów, Szczerców, and Turów opencast mines. Its macroscopically estimated amount is the largest, attaining up to 30% of seam thickness, in the case of ŠLS-3 in the Turów opencast mine (Fig. 5E; Wagner, 1996; Kasiński, 2000, Kasiński et al., 2010). However, the total geological resources of Polish bitumen-rich lignites were calculated to be merely about 0.64 Mt at the end of 2014 (Szuflicki et al., 2015).

3.5. Fusitic lignite association (FL)

3.5.1. Description
The fusitic lignite is characteristically black or silky black, friable with needle-shaped particles, and highly porous (Wolf, 1988; Markič & Sachsenhofer, 1997; Kolcon & Sachsenhofer, 1999). Therefore, its bulk density is very low (average 0.3–0.6 g/cm³), allowing it to be easily distinguished from other lithotypes (average > 1.1–1.3 g/cm³). Generally, the fusitic lignite is rare within Polish lignite seams, an exception being the Sieniawa opencast mine, where it forms a few well-developed horizons that are up to 5–10 cm thick (Fig. 6A–C). The fusitic lignite creates lenses more frequently than it does continuous strata. Its structure may be different; however, it is predominantly nodular (Fig. 6C, D).

3.5.2. Interpretation
In essence, the fusitic lignite is a fossil charcoal that is produced by the combustion of plant matter by both exogenous and endogenous processes. In the former case, the dry parts of the mires may be set alight by lightning or volcanic eruptions. In the latter, spontaneous combustion of a lignite seam can be caused by bacterial activity or pyrite oxidation. In any case, the presence of fusitic lignite in the seam is evidence of natural fires in a mire or within the lignite seam (e.g., Brzyski, 1981; Kwiecińska & Wagner, 1997; Ticleanu et al., 1999; Majewski, 2002).

3.6. Weathered lignite association (WL)

3.6.1. Description
The distinction of this lithotype association is sometimes very useful in the field; which is why it is added to the present paper (compare Widera, 2012). It is easy to distinguish from other lithotype associations on account of its colour, which is mostly yellowish brown to reddish brown (Fig. 7A). Moreover, it forms a continuous layer over a stretch of at least several tens of metres and a thickness of up to 50 cm. This lithotype association can be characterised by various structures that are typical of all Polish lignites (Fig. 7; Table 1); however, only a folded structure has been documented (Fig. 7A).

3.6.2. Interpretation
The weathered lignite results from moisture deficiency on the mire surface. The upper layers of fresh peat were then oxidized under arid conditions, when the groundwater table was relatively low and/or long-lasting droughts occurred. There-
Marek Widera

4. General lithotype composition of lignite seam exploited

Among the currently exploited Polish lignite seams, the highest content of xylites is within MPLS-1 in all opencasts of the Konin and Adamów lignite mines. It has a xylite content of about 8% (by volume, i.e. roughly equal to the 2-dimensional; compare Fig. 3), including almost 4% fibrous xylites. Conversely, ŚLS-3 and LLS-2 are characterised by a total xylite content of less than 7% and a fibrous xylite content of less than 1% (Bielowicz, 2013). However, the xylite content in some sections may range from 40% in LLS-2 in the Turów opencast to 55% in ŚLS-3 and LLS-2 (which, together, create the so-called main coal seam) in the Belchatów open-cast (Wagner, 1996).

In the case of lithotype composition, it should be stated that the lignite seams exploited differ significantly from each other (Table 2). In the Turów opencast, for instance, ŚLS-3 consists of 33.5% xylodetritic lignite (XDL) and 27.2% detritic lignite (DL). This is in contrast to ŚLS-3 and LLS-2 from the Belchatów mine, where the above lithotype associations together contain 34.2% detroxylitic lignite (DXL) and 30.4% xylodetritic lignite (XDL). However, the sum of these two lithotype associations (that is, DXL and XDL), is very similar in the case of lignites from the Bełchatów and Sieniawa mines. MPLS-1 is extracted only at the Konin and Adamów mines. Moreover, MPLS-1 differs between the two mines. It consists of 42% DL at the Konin Lignite Mine, but 85% XDL in the Adamów Lignite Mine (Table 2).

The contribution of other lithotype associations within the lignites examined also appears of interest, although it varies greatly. First, the relatively high percentage (9.8–11.8%) is a share of bitumen-rich (bituminiferous) lignite in seams ŚLS-3 and LLS-2 exploited in the Belchatów and Turów opencasts. Second, weathered (WL) and fusitic (FL) lignites are also present in the Sieniawa opencast. Macroscopically, these lithotype associations are easily noticeable there, because they comprise 13% and less than 1% of the thickness of LLS-2, respectively (compare Table 2 and Figs 6A–C, 7A).

Finally, the size and degree of gelification of the Polish lignite seams require a few words here. The

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Fig. 6. Fusitic lignites with a folded and nodular structure. A, B, C – Examples of fusitic lignites from the 2nd Lusatian Lignite Seam (LLS-2) in the Sieniawa opencast mine, Sieniawa Lignite Mine; D – Fusitic lignite from the 1st Middle-Polish Lignite Seam (MPLS-1) in the Drzewce opencast mine, Konin Lignite Mine. For location see Figure 1, for stratigraphical position of the lignite seams see Figure 2 and for explanation of lignite lithotype associations codes see Table 1.
An overview of lithotype associations of Miocene lignite seams exploited in Poland

Fig. 7. Lignite lithotype associations with a deformed structure of various origin. A – Glaciotectonically folded lignites within the 2nd Lusatian Lignite Seam (LLS-2) in the Sieniawa opencast mine, Sieniawa Lignite Mine; B – Fractured detritic and xylodetritic lignites – the fractures represent face and butt cleats within the 1st Middle-Polish Lignite Seam (MPLS-1) in the Józwin IIB opencast mine, Konin Lignite Mine; C – Lignites folded by the weight of huge, peat-forming trees within the 1st Middle-Polish Lignite Seam (MPLS-1) in the Drzewce opencast mine, Konin Lignite Mine; D – Faulted lignites within the 2nd Lusatian Lignite Seam (LLS-2) in the Sieniawa opencast mine, Sieniawa Lignite Mine; E – Faulted lignites within the 3rd Ścinawa Lignite Seam (ŚLS-3) in the Bełchatów opencast mine, Bełchatów Lignite Mine. For location see Figure 1, for stratigraphical position of the lignite seams see Figure 2 and for explanation of lignite lithotype associations codes see Table 1.

Table 2. Average composition of lithotype associations of currently exploited Polish lignites (after various authors).
For location of lignite mines see Figure 1, for explanation of lignite lithotype associations codes see Table 1 and for explanation of lignite seam codes see the text.

<table>
<thead>
<tr>
<th>Lithotype association</th>
<th>Turów Lignite Mine¹</th>
<th>Belchatów Lignite Mine²</th>
<th>Sieniawa Lignite Mine²</th>
<th>Konin Lignite Mine³</th>
<th>Adamów Lignite Mine⁴</th>
</tr>
</thead>
<tbody>
<tr>
<td></td>
<td>ŚLS-3</td>
<td>LLS-2</td>
<td>ŚL-3 &amp; LLS-2</td>
<td>LLS-2</td>
<td>MPLS-1</td>
</tr>
<tr>
<td>XL</td>
<td>6.3%</td>
<td>7.6%</td>
<td>3.0%</td>
<td>3%</td>
<td>10%</td>
</tr>
<tr>
<td>DL</td>
<td>27.2%</td>
<td>4.6%</td>
<td>20.6%</td>
<td>15%</td>
<td>42%</td>
</tr>
<tr>
<td>DXL</td>
<td>7.0%</td>
<td>22.6%</td>
<td>34.2%</td>
<td>28%</td>
<td>35%</td>
</tr>
<tr>
<td>XDL</td>
<td>33.5%</td>
<td>22.7%</td>
<td>30.4%</td>
<td>40%</td>
<td>10%</td>
</tr>
<tr>
<td>BL</td>
<td>11.8%</td>
<td>11.5%</td>
<td>9.8%</td>
<td>no data</td>
<td>no data</td>
</tr>
<tr>
<td>FL</td>
<td>present</td>
<td>present</td>
<td>present</td>
<td>&lt; 1%</td>
<td>&lt; 0.1%</td>
</tr>
<tr>
<td>WL</td>
<td>present</td>
<td>present</td>
<td>present</td>
<td>13%</td>
<td>no data</td>
</tr>
</tbody>
</table>

lignites from the Turów and Belchatów mines generally are more strongly gelified than those from the Konin and Adamów mines (Fabiańska & Kurkiewicz, 2013). This simply means that ŚLS-3 and LLS-2 are much more gelified than MPLS-1. Furthermore, the degree of gelification increases from the bottom of ŚLS-3 to the roof of LLS-2 in the Belchatów, Szczerków, and Turów opencasts. Conversely, LLS-2 in the Sieniawa opencast is weakly gelified. In this case, it is more closely similar to MPLS-1 from the Konin and Adamów mines than LLS-2 and/or ŚLS-3 from the Belchatów and Turów mines.

5. Discussion

5.1. Lignite lithotype associations vs lignite seam stratigraphy

The macroscopically estimated composition of the lithotype associations is very different for ŚLS-3, LLS-2 and MPLS-1, which are currently mined in Poland. There is no close relationship between the lithotype associations examined and the stratigraphical position of the lignite seams. Generally, xylodetritic (XDL) and detoxylic (DXL) lignites, which together constitute 40.5–87% of the seam thickness, predominate among Polish lignites (Table 2). The average content of fibrous xylites differs strongly in various stratigraphically lignite seams. Thus, the youngest seam, MPLS-1, includes at least four times more fibrous xylites than the older ones, LLS-2 and ŚLS-3 (Bielowicz, 2013). Most likely, the lithotype composition is rather the result of the original plant assemblages than a reflection of age and/or stratigraphical position of the examined Miocene-aged lignite seam.

Predominantly, the presence of fusitic and weathered lithotype associations within the lignite seams is evidence of the dry environmental conditions during development of the mire. Both of these lithotype associations are definitely the commonest in the Sieniawa mine (compare Table 2 and Figs 6A–C, 7A). The bitumen-rich lignite cannot be also related to the stratigraphy of the lignite seams. This opinion confirms the lack of this lithotype association within LLS-2 at the Sieniawa mine, although its content within LLS-2 and ŚLS-3 in the Belchatów and Turów mines is 9.8–11.8% (Table 2; Wagner, 1996).

Similar to the bitumen-rich lignite, the gelified structure does not exist or is negligible in lignites from the Sieniawa mine. Conversely, it is very characteristic of coeval and older lignites from the Belchatów and Turów mines (compare Fig. 5A–D). Therefore, the stratigraphical position of the lignite seams studied does not directly determine its degree of gelification.

5.2. Lignite lithotype associations vs lignite seam location

The relationship between the peat-forming environments and the resultant lignite lithotype associations generally does not depend on stratigraphy (as discussed above) or location of the exploited lignite seams. It is possible to indicate such a close dependence in at least two cases that are related to the bitumen-rich and gelified lignites, which are present in the Belchatów and Turów mines but not in the Sieniawa mine. Because of their practical importance and scientific specificity they need to be discussed here.

In the case of the bitumen-rich lignites (BL), the similarities of those from the Belchatów and Turów mines are easily noted. Moreover, differences between the Belchatów and Turów bitumen-rich lignites and lignites from the Sieniawa mine are very large (compare Table 2). As stated above, the bitumen-rich lignite formed in the shallow water of lakes that covered the mire surface or existed in close proximity to the mire (compare Traverse, 1988; Wagner, 1996; Kasirski et al., 2010). Therefore, the discussed similarities and differences can be explained by various palaeogeographies of the lignite-bearing areas in Poland during the formation of ŚLS-3 and LLS-2. At that time, the mires developed in the proximity of relatively long-lasting lakes in the vicinity of Belchatów and Turów lignite deposits. On the other hand, the area of the Sieniawa lignite deposit was located at some distance from such lakes during deposition of LLS-2.

Strong gelification of lignites from the Belchatów and Turów mines corresponds strictly to their location. It should be noted that the lignite seams examined (ŚLS-3 and LLS-2) are located in tectonic grabens, which are > 500 and > 250 m deep, respectively (e.g., Kasirski, 2000; Kasirski et al., 2010; Widera & Haluszczak, 2011; Widera, 2016a, 2016b). However, the cause of such a high degree of gelification is different in each case under consideration.

ŚLS-3 and LLS-2 (the main coal seam) from the Belchatów mine are strongly gelified due to the very high calcium (Ca) content in some parts of the mire. The source of CaCO₃ lay in carbonate Jurassic and Cretaceous rocks that are elevated around the lignite deposit. Thus, it seems plausible that
CaCO₃ was supplied to the graben from its margins (e.g. Wagner, 1996; Wagner et al., 2000; Mastej et al., 2015). This explanation is consistent with the opinion of other researchers, who assumed a close relationship between mire alkalinity and degree of lignite gelification (e.g., Diessel, 1992; Markič & Sachsenhofer, 1997; Bechtel et al., 2003).

The Turów lignite deposit is located in the northeastern segment of the Eger (Ohře) Graben (e.g., Kasiński, 2000; Kasiński et al., 2010; Widera, 2016b). It is surrounded and underlain by Precambrian and Cenozoic volcanic rocks. On the other hand, the degree of gelification is considerably stronger at the bottom of ŚLS-3 than at the roof of LLS-2 as mentioned above. Therefore, the increased degree of geothermal energy in southwest Poland, especially during Early–Middle Miocene volcanic activity, is most likely responsible for the very strong gelification of lignites in the Turów mine (compare Fig. 5A, 5B).

Finally, it is noteworthy that the bitumen-rich (bituminiferous) lignites and gelified lignites have advantages and disadvantages during the utilisation process. Their presence within the studied seams significantly increases the lignite’s calorific value, which is closely associated with a higher rank of coal/lignite. Some parts of the seams with high quantities of the bitumen-rich and gelified lignites are characterised by limited usefulness, for instance, in the process of underground gasification. On the other hand, a relatively large amount of fibrous xylites is also undesirable in lignite mining, combustion and possible gasification. This is because fibrous xylites make it difficult for chain excavators to work, are characterised by lower calorific value than other lithotypes, and are hard to grind in the process of lignite combustion and gasification (e.g., Wagner, 1996; Kwiecińska & Wagner, 1997, 2001; Bielowicz, 2013; Bielowicz & Kasiński, 2014).

6. Conclusions

Polish lignites are highly diversified lithologically and petrologically. The currently exploited lignite seams represent a wide range of lithotypes with characteristic textural and structural features. Each of these lithotype associations corresponds or may be related to specific environmental conditions of peat deposition. Such peat-forming environments and the resultant lignite lithotypes are known from all humic ortholignites of low-rank coal deposits around the world.

In most cases there was no correlation between the lithotype associations’ composition and the stratigraphy of the investigated lignite seams or between the lithotype associations’ composition and the location of these seams. Such a close relationship was found only with respect to the bitumen-rich and gelified lignites from the Belchatów and Turów mines. In the case of the former, this is explained by the palaeogeographical conditions. The increased degree of gelification within the lignite seams in these two mines is linked to mire alkalinity in the Belchatów Lignite Mine and to the degree of geothermal energy in the Turów Lignite Mine.

In summary, the lithotype associations’ composition is important for mining activities and multi-use of lignite. An accurate description of the currently exploited lignite seams, including the determination of lithotypes, may not only be of great importance for scientific knowledge but also useful for present or future utilisation of lignite in Poland. Therefore, further studies of lignite lithotypes are encouraged.

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